

**REVISED SEISMICITY OF THE GRAND BANKS
AND OFFSHORE NEWFOUNDLAND**

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ABSTRACT

A comprehensive review of the catalogued seismicity of the Grand Banks of Newfoundland (1922–1994) has revised or added earthquakes in four time periods: 1922–1970, the two largest earthquakes ($M \approx 5$) relocated to Baffin Island (1922) and the northern Labrador Sea (1962); 1977–1982, added four new earthquakes; 1969–1984, revised the locations and magnitudes for five earthquakes; 1984–1994, added 15 earthquakes. Many earthquakes appear to be spatially associated with Mesozoic extensional structures that trend northeastwards across the Grand Banks. A cluster of six post-1970 earthquakes lies 40–80 km northwest of the Hibernia oilfield and represents the most seismically active region of the Grand Banks.

RÉSUMÉ

Un examen approfondi de la sismicité cataloguée antérieurement à la région des Grands Bancs de Terre-Neuve (1922–1994) a révélé quelques révisions et quelques séismes nouveaux dans les quatre périodes suivantes: 1922–1970, relocalisation des deux séismes les plus importants ($M \approx 5$) à l'île Baffin (1922) et dans le nord de la mer du Labrador (1962); 1977–1982, quatre séismes nouveaux; 1969–1984, révisions aux localisations et aux magnitudes de cinq séismes; 1984–1994, ajout de 15 séismes. Plusieurs séismes semblent associés spatialement à des structures extensionnelles du Mésozoïque qui tendent vers le nord-est à travers les Grands Bancs. Une grappe de six séismes postérieurs à 1970 se trouve de 40 à 80 km au nord-ouest du champ pétrolifère Hibernia et constitue la région la plus active des Grands Bancs.

INTRODUCTION

This open-file is the third of a series documenting our current knowledge of earthquakes off the southeastern Canadian continental margin (Fig. 1). The other reports in the series include: Adams and Staveley (1985) for the island of Newfoundland, and Adams and Simmons (1991) for the Labrador Sea, and are expected to include: Adams (in prep.) for the Laurentian Slope seismic zone, Adams (in prep.) for the Laurentian Channel and Fan, and Wahlström and Adams (in prep.) for the Scotian Margin. The current report covers a part of the region defined by the Grand Banks Basin Atlas sheet (a Frontier Geoscience project of the Geological Survey of Canada), and revises the epicentres and magnitudes of earthquakes on the Grand Banks and adjacent parts of the offshore.

As a side note, it should be noted that the "Grand Banks" earthquake of 1929, at magnitude 7.2 the largest on the southeastern Canadian continental margin, occurred in the Laurentian Channel (44.5°N 56°W), south of Newfoundland, and not on the Grand Banks proper. Thus it is not discussed in the present report.

Besides the more long-term scientific aspects, a recent interest in the seismicity of the Atlantic Margin has arisen from the need to produce seismic design levels (e.g. Foo and Crouse, 1986) for the Hibernia platform, near 46.7°N 48.9°W, east of St. John's, and to refine estimates of seismic hazard to future offshore hydrocarbon production platforms such as may be emplaced off Sable Island, east of Halifax, or on the Grand Banks (Keen et al., 1990). Earlier hazard estimates (Basham et al., 1983; Adams, 1986) recognized the poorly determined nature of the offshore seismicity and the need to improve seismograph coverage to detect more of the smaller offshore earthquakes and to relocate the older earthquakes.

The current project was begun in 1982 by Adams, and many of the original earthquake phase arrivals were re-read during R. Wahlström's study period in Ottawa in October 1984 - October 1985. Preliminary epicentres and magnitudes for the older earthquakes were determined at that time. Although only informal, these results have been available for use in seismic hazard analyses and have appeared in various reviews (e.g., Basham and Adams, 1983; Adams, 1986; Adams and Basham, 1989); it is our intention here to finally report them

in a formal manner, and to update the results to take account of more recent earthquakes, 1985-1994.

HISTORY OF EARTHQUAKE MONITORING

A brief, and probably still incomplete, history of earthquakes in Newfoundland is given by Adams and Staveley (1985). Prior to the advent of instrumental monitoring of earthquakes in Atlantic Canada, all earthquakes were detected by their felt effects. All of the earthquakes felt on the island of Newfoundland documented by Adams and Staveley (1985) appear to pertain to small, onshore earthquakes and not to large offshore earthquakes. By contrast, some earthquakes felt in coastal Labrador are likely to have been large offshore events (Basham and Adams, 1983; Adams and Staveley, 1985), though A. Stevens (pers. comm., 1991) considers they might have been local earthquakes along the coast. For example, an event felt widely (but not strongly in any one place) in Nova Scotia and New Brunswick in 1882 might have been an offshore earthquake (Ruffman and Peterson, 1988, p. 284). Another feature that might signify an offshore earthquake is the occurrence of a tsunami (e.g., Adams and Staveley, 1985; Ruffman et al., 1991); however no earthquake has yet been convincingly confirmed from the numerous reports of tsunami-like waves.

Seismograph monitoring of the Grand Banks has improved with time, so that small earthquakes that can currently be detected would have passed unnoticed in previous years. The following brief history is adapted in part from Stevens (1980). A seismograph was operational in Halifax (HAL) from 1915, but it operated at a low magnification (first a Bosch and then a Mainka instrument, with magnifications of 100 - 150x) and was able to detect only the largest earthquakes. The installation of the Wood-Anderson seismographs (peak magnification at 1 second = 2700x) at Seven Falls (SFA) and Shawinigan Falls (SHF) in 1927 improved the monitoring, as did the Benioff seismograph ($\approx 8,000x$ for 1937-1939 and $\approx 80,000x$ subsequently) at Ottawa in 1937 and the Willmore seismograph ($\approx 30,000x$) at Halifax in 1952.

All of these seismographs lay distant from and southwest of the Grand Banks, and it was not until 6-component Canadian Standard Seismographs were opened at Schefferville (SCH) in July 1962 and St. John's (STJ) in June 1964 that coverage improved significantly. STJ, while it was well-located with respect to Grand Banks earthquakes, operated at a magnification of only 35,000x for the short-period instruments because of the high microseismic and cultural noise of the site.

A further improvement in monitoring occurred when Memorial University of Newfoundland started to operate short-period vertical component seismographs in St. John's (in the same vault as STJ, but herein named 'MUNF' to distinguish it for the period when the STJ standard station was also in operation) in February 1975 and Corner Brook (CBK) in October 1976. These instruments operated at a higher gain ($\approx 100,000x$) than STJ because of their restricted low frequency response and proved very successful at detecting smaller offshore earthquakes. However, they were not always operated to modern standards with respect to timing, calibration, or record annotation, and the records were not routinely read with the other Canadian seismograms. In June 1983, Adams retroactively read the earlier seismograms, and the subsequent seismograms have been forwarded to Ottawa and have been systematically read by Adams as part of the national seismograph network.

The most recent changes affecting the earthquakes in this report are as follows. STJ was closed in March 1989, and the MUNF seismometer was relocated from a vault on the main campus to nearby Mount Scio and was renamed to STJN in September 1991. STJN has been operated under contract from Memorial University to the Geological Survey of Canada since April 1989, and CBK was operated under contract from April 1989 to March 1994. GSC's new digital seismograph at Deer Lake (DRLN) opened in December 1993 and replaced the function of CBK.

Although coverage has improved with time, it is still poor with respect to the rest of Canada. As will be seen (e.g., Fig. 4), very few earthquakes of magnitude less than 4 are detected more than 300 km distant from St. John's. The few larger earthquakes that are located (or only detected, see Table 2) must therefore be indicative of many smaller earthquakes that pass unnoticed in the same active areas. The last earthquakes mapped in this report occurred in 1992. There were no earthquakes located in the study area in 1993; the three that occurred in 1994 and one in early 1995 are added only to the Appendix and

the "Results" section as a "note in proof", since the report was substantially completed in July 1992, even though its issue was delayed until March 1995.

PROCEDURES

Choice of events studied

Figure 1 delimits the area of the southeastern offshore covered by the Basin Atlas mapsheets and the sub-region treated in this report. The earthquakes studied in this paper represent: all known earthquakes thought to lie in the Grand Banks region according to the Canadian Earthquake Epicentre File (CEEF) in October 1985 (being the previous state of knowledge, see Fig. 2 and Table 1); all subsequent earthquakes located by current methods; and some older earthquakes (not in the CEEF) found by searching the older Memorial University seismograms.

In all, 23 pre-1993 earthquakes are studied in this paper: the two earliest are moved away from the area (having been grossly mislocated), 4 are new events added from the Memorial seismograms, and 17 are earthquakes which have either been relocated (pre-1985 events) or new events located during this project. In addition a number of unlocatable events were detected, usually only on STJ or MUNF, but occasionally on CBK alone if the high-gain MUN station in St. John's was not operating (Table 2). These earthquakes were too small to register on other seismographs and are thus probably offshore.

Determination of Epicentres

For all the earthquakes studied here the epicentres and focal depths have been determined with the standard crustal model and location program used by the Geophysics Division. This model is a good average for the Canadian Shield and is used for offshore earthquakes in the absence of sufficient information to refine the average velocity structure. The crustal model has a 36-km-thick crust of P-velocity 6.2 km/s overlying a mantle of P-velocity 8.2 km/s. S-wave velocities are 3.62 km/s (revised from the previous crustal 'Lg' velocity of 3.57 km/s by Connors and Adams, 1988) and 4.70 km/s, respectively. It would

not be surprising if the use of this simple structure leads to some systematic bias in the epicentres (likely up to a few tens of kilometres, but possibly even larger), because with the distribution of land-based seismographs (all to the east of the earthquakes) the epicentres move to minimize the residuals created by an inappropriate crustal model. However, it is impossible to be sure because there have been no earthquakes large enough to give a reliable teleseismic epicentre for comparison, and no earthquakes located both by ocean-bottom seismometers and the seismograph network (contrast the situation on the west coast of Canada, e.g. Wahlström and Rogers, 1993; Wahlström et al., 1990). For these reasons the reader should be warned that the apparent precision of the results (e.g. the errors on Table 3 or in the the Appendix "pikfile" solution lines) is misleading. Furthermore, because of the distribution of stations, the true epicentres are often more uncertain in latitude than longitude, even where this is not indicated by the errors in Table 3.

For each earthquake, all existing Canadian seismograms were read for phase identification, phase onset times (Pn, Pg, Sn, Lg), maximum amplitudes and corresponding periods (Sn or Lg), and P-wave polarity (Pn or Pg). Many of the records for the earlier earthquakes were read by Wahlström at Lamont-Doherty Geological Observatory, New York, where the older Canadian seismograms were stored until the mid-80's. Adams subsequently re-read selected phases in Ottawa. Phases from the Greenland stations were added from Danish bulletins when available. Timing was uncertain on the Memorial seismogram until about 1983; notes in the Appendix indicate where an arbitrary time correction was applied in order to use the S-P information.

The routine determination of epicentres by the Geophysics Division uses all available phases except those that obviously misfit. Hypocentral depths are often assigned to 18 km, being simply the mid-crustal depth according to the model. Epicentres in the present paper were recalculated from the earthquake phase arrival times according to the following guidelines:

- 1) STJ and MUNF were co-located in the same vault. Therefore where individual phases were read on both instruments, only one was used in the location, the other being zero-weighted ('X-ed' out). Usually it was the Pn from MUNF (its higher gain gave a sharper arrival) and the Sn from STJ (the Sn could be read best from the horizontal component; MUNF was vertical component only) that were used.

- 2) Pg and Sg arrivals for stations within about 100 km were fit well, even at the expense of larger residuals on the many distant stations (affects only events #69, #94b, #94c, and #95).
- 3) Pn and some Sn arrivals for the closest stations were fit well, at the expense the fit of arrivals at more distant stations.
- 4) Lg phases at distances beyond about 1000 km were given lesser weight because of their less sharp onsets and uncertainty in the Lg velocity.
- 5) In a few cases, a single station was chosen to replace a cluster of stations at a similar azimuth and distance to avoid undue bias.

The results of this data selection can be seen by examining the residuals for the earthquake solutions in the Appendix. While such recomputations of standard epicentres – which occasionally revise the routinely-determined epicentres by many kilometres – do not guarantee that an **accurate** epicentre has been found, we believe that they represent an improvement on the routinely determined epicentres. The epicentres are listed in Table 3 and plotted in Figure 3.

Focal depth

It is in general difficult to compute focal depth for small local earthquakes in the absence of a very dense local seismic network and a good knowledge of the crustal velocities. For large ($M > 5$) earthquakes that are recorded at teleseismic distances (> 2000 km) it is possible to compare arrivals of different waves to determine the depth of the earthquake, e.g. the delay between the direct, downward P-wave and the wave that propagates upwards before being reflected down at the earth's surface as P (pP and sP). All the earthquakes studied here are too small for this method.

At close epicentral distances (conventionally, within twice the focal depth of the earthquake being studied), the depth can be computed by minimizing the combined residuals on the upward- and downward-propagating rays; however, the earthquakes in this report lie offshore and have poor seismograph coverage at small distances. Consequently all but one earthquake in this report are placed at the mid-crustal depth (18 km) of the standard model.

Determination of magnitudes

Magnitudes are determined according to current GSC practice, namely: a m_b or M_s from teleseismic P amplitudes is preferred for the larger events (say $M > 5$) if it is available from USGS or ISS; a m_{bLg} (also described as a m_N) determined from the Lg-wave amplitude is preferred where the Lg has propagated normally and not been severely attenuated; and finally a M_L determined from the Sn amplitude is used as a last resort. Supplementary magnitudes (e.g. m_b where a m_{bLg} is preferred) are retained in the CEEF. The M_L calculation with Sn amplitudes uses California attenuation and overestimates magnitudes from stations involving long paths across the high-Q Canadian Shield. Results reported by Adams and Simmons (1991), who compared teleseismic m_b with M_L for nine Labrador Sea earthquakes, suggest that the M_L 's for the magnitude range 4.5 to 5.4 might be higher than m_b by 0.4 magnitude units. However, their results are too poorly based to suggest that the M_L magnitudes reported here should be uniformly reduced by that estimate of bias.

Determination of polarities

Where possible, polarities were read from the first P (Pg or Pn) on all available records. These produced a sparse set of data, often only the polarity on the closest station being readable. Polarities were assigned either full or half weight. The polarities were read by both authors and readings on which we differed were either not used or a consensus was reached and the polarity was used at half-weight. Full weight polarities are impulsive and unambiguous. Half weight polarities are emergent, less strong, or may occur on noisy records.

RESULTS

The results of the relocations are summarized in Tables 3 and 4, and as a map of the revised seismicity in Figure 3. Details of the relocated solutions are in the Appendix, which also contains detailed comments on the phases and solutions. Figure 4 shows a map of the relocation vectors. The earthquakes are mapped at their revised locations

and the tails point back to their old epicentre. Figure 5 shows the new epicentres with diamonds showing the computed precision of the epicentre (large diamonds represent poorly determined epicentres), Figure 6 shows the detail of the area east of St John's, and Figure 7 shows the detail of the cluster of earthquakes near Hibernia.

In the following discussion, and in Tables 3 and 4 and the Appendix, earthquakes are denoted by their date, e.g., #89b, which is the second earthquake in 1989.

#22 was originally located by the International Seismological Summary (ISS) at 50°N , 50°W based on six stations, only OTT being in Canada. Although SAS was operating, the records were apparently not saved, and the records for VIC have still not been located. The OTT records were not found, and the microfilm of the Bosch and Milne-Shaw records show no clear signal. However, we did find the HAL Mainka record (one horizontal component). This had a time correction of -45 seconds, and gave clear phases $P=06\text{h}34\text{m}33\text{s}$, amplitude 0.10mm at a period of 2.6 s; $S=06\text{h}39\text{m}05\text{s}$; and surface wave maximum at 06h42m amplitude 0.15mm at a period of 4.0 s. This indicates a distance of 26.5° (irrespective of the exact time correction; that P and S are the recorded phases is confirmed by later arriving surface waves), which is more than twice the distance of HAL from the assigned ISS epicentre (the HAL readings were not included in the ISS solution).

The distance from HAL is also consistent with the recorded surface wave maximum at 06h42m. Using the arrival times given by ISS and a revised phase interpretation we find origin time 062855, OTTAWA distance= 25° (S and RM; the P misfits by about 1 min, which may be a 1 minute reading or reporting blunder because accurate timing was not a problem at the Dominion Observatory), CHICAGO distance= 28.5° (RM), WASHINGTON distance= 32° (RM), and ESKDALEMUIR distance= 32° (RM). 'RM' denotes the Rayleigh wave maximum, the time of which can be used for crude locations in the absence of phase onset arrival times (At Uppsala a set of RM travel time tables prepared by Båth are used). These readings give an approximate epicentre at 70°N , 75°W (Baffin Island). For this solution, the two other ISS readings give distances about 4° too long if the recorded phase is interpreted as RM for this solution (BIDSTON distance= 38.5° and DE BILT distance= 41°). The

Uppsala Weichert records were checked, but there was absolutely no trace of the earthquake. Thus, that two of the three European stations with data in the ISS are off by a few minutes should not be discouraging: the signals must be very weak.

Dr. Anne Stevens has suggested that the HAL, OTT, and WAS phases would fit neatly as Pn, Lg; Pn, Sn; and Lg respectively and give an epicentre near 59N 53W, off southwestern Greenland. We can not agree with this location, for: 1) if the epicentre is in eastern Labrador Sea, then the second phase at HAL is not Lg, because Lg does not propagate across this oceanic region (the same goes for the WAS reading); and 2) the clear 3rd phase at HAL is not explained. The "reasonable fit" of these three stations is worth little compared with our new solution which satisfies practically all stations.

Our solution, with identified and accurately timed phases at HAL, giving a reliable distance, and several fitting LR maximum arrivals for a Baffin Island location, is no doubt the best from the documented data. The greatest "obstacle" is the misfitting OTT P arrival and that this cannot be checked since the records are lost. Our solution is both the best and is very rough (see Discussion below).

Smith (1962), gave a fictitious intensity of MM VI-VII to this earthquake, which was not reported felt, and from this intensity the current M_L 5.3 in the CEEF was derived. From the horizontal component at HAL (orientation unknown) we read a trace amplitude of 0.15 mm at a period of 4.0 s. Assuming i) a magnification of 150 (the maximum magnification was 100-150 according to Stevens, 1980) for the found record, and ii) zero amplitude for the missing orthogonal component, we compute $M_s(\text{Prague-Moscow})=5.1$. Due to the two assumptions this is the minimum value.

In conclusion #22 is a magnitude 5 earthquake in Baffin Island (about 2850 km from its ISS epicentre) and not a Grand Banks earthquake.

#62 was originally located east of the Strait of Belle Isle, and prior to revision was a single large earthquake in an otherwise aseismic area. Both the original and revised locations are based on only three phases from two stations, a P and an S on SCH and one phase on HAL, which Smith (annotations on Dominion Observatory epicentre cards) took to be Sn. A inland epicentre west of Schefferville was ruled out by Smith because the earthquake was not detected on southern stations such as SFA or

MNT, so he located it on the southern Labrador Shelf. After re-examining the HAL seismogram and confirming that only one phase was recorded, we have re-interpreted Smith's Sn reading as the Pn phase. This, together with the absence of Lg energy at Schefferville, is consistent with an earthquake in the oceanic crust of the Labrador Ridge. A slightly larger Labrador Ridge event, 621202, close to the revised #62 epicentre, is well-located, and is also found to lack a Sn reading for HAL (Adams and Simmons, 1991). HAL and SCH were at similar distances from the old #62 epicentre, but the SCH record had a 30 mm amplitude, while HAL signal was "very small"; this suggests the earthquake was much further from HAL, as we now locate it. The epicentre of #62 has thus been revised by over 1000 km.

#69 occurred very close to St. John's, but was not reported in the newspapers as being felt (Adams and Staveley, 1985). The epicentre is revised 17 km to the SE.

#70 is a relatively large earthquake, M_L 4.3, north of the Grand Banks. It is located using only stations STJ and SCH, but the absence of phases on HAL excludes the southern alternative epicentre.

#71 is now the largest earthquake on the Grand Banks at m_b 4.8. It lies within 65 km of the Hibernia Oilfield, and has been moved 32 km SE of the CEEF epicentre. The maximum amplitude readings on Canadian seismographs for this earthquake represent Lg energy, and the m_N 4.8 so calculated is similar to the teleseismic m_b from the ISC. However, m_N is the preferred magnitude scale since it best represents the frequencies of engineering interest when this earthquake is used in eastern Canadian seismic hazard calculations.

#76 is confirmed as being located north of the Grand Banks, the revised epicentre being essentially the same as that in the CEEF.

#77 is a new earthquake found by searching the Memorial University seismograms, but was too small to be confirmed by GSC seismograph records. From the absence of phases at Halifax, the earthquake lies to the northeast of St. John's on the northern Grand Banks, rather than south of St. John's.

#78 is a new earthquake found by searching the Memorial University seismograms, and confirmed by GSC seismograph records. It lies far to the north of the Grand Banks, close to the similar-sized #70.

#81 is a new earthquake found by searching the Memorial University seismograms, and confirmed by GSC seismograph records. It is one of the most easterly of the Grand Banks earthquakes, but some Lg energy did propagate to STJ.

#82 is a new earthquake found by searching the Memorial University seismograms, and confirmed by GSC seismograph records. It is the first known earthquake from the Newfoundland Ridge - Tail of the Bank area. It is magnitude $M_L 4.0$, a relatively large earthquake to have been missed during the routine compilation of seismicity (it occurred during the Miramichi aftershock sequence and may simply have been overlooked). However, it is probably close to the detection threshold for this remote area, and should be taken as indicative of seismic activity along the southeastern margin of the Grand Banks that is currently undetectable.

#84a is a moderate earthquake on the northern Grand Banks, notable for having a $M_L 2.6$ #84b aftershock two and a half hours later. The mainshock epicentre is relocated 17 km to the SE of the CEEF epicentre. The aftershock was entered into the CEEF as a note, but here is proposed to have its own entry (#84b). The STJ seismograms for both mainshock and aftershock show significant Lg energy, indicating a source in continental rather than thinned continental or oceanic crust.

#85 is an earthquake found first by searching the Memorial University seismograms, and then confirmed by GSC seismograph records. By contrast to #84a and #84b, the Lg wave did not propagate efficiently to STJ, suggesting the earthquake occurred in oceanic or thinned continental crust. #85 lies farther offshore from St. John's than #84, and may reflect crustal thinning towards the edge of the Grand Banks.

#88a is similar to #85, occurs in the same general area, and also lacks Lg energy at STJ. The current epicentre is 10 km to the SW of the CEEF epicentre.

#88b occurred close (within 20 km) to another large earthquake (#71) in the vicinity of the Hibernia Oilfield. Lg-wave energy was not recorded at STJ, even though the STJ seismogram of #71 showed an Lg-wave.

#89a is a large, $M > 5$, earthquake near the Mid-Atlantic Ridge, which is included for completeness. Many earthquakes occur on the ridge to the east of the map area, and some of the larger ones give Pn and sometimes Sn energy on Atlantic Canada seismographs. These earthquakes are best located using world-wide data (we choose the ISC solution

instead of computing a GSC solution), and have no seismic hazard implications for Canada, because they are too distant to cause significant ground shaking and these mid-Atlantic Ridge earthquakes are not known to trigger tsunamis.

#89b is an earthquake on the Grand Banks. It appears to be close to the Hibernia Oilfield and is the largest earthquake in the vicinity since 1971.

#90 is only the second known earthquake from the southernmost Grand Banks. It is considerably larger (M_L 4.7) than #82, which is nearby. The m_b determined by the ISC from 8 stations is 4.3, consistent with the anticipated overestimation of magnitude by M_L .

#91 is an earthquake on the shelf east of St. John's. It is clearly closer to St. John's than the cluster of events (#71, #88b, #89b, #92a,b,d) that lie about 80 km farther east.

#92a is the third magnitude 4 earthquake to occur in a cluster about 270 km east of St. John's. This earthquake was well-recorded throughout eastern Canada, with Lg-wave energy being seen as far as Alberta. Relative arrivals at St. John's and Halifax suggest a similar location to the much larger 1971 earthquake (see Discussion).

#92b is a further earthquake in the cluster about 270 km east of St. John, the second of three in 1992.

#92c is the first earthquake known from the eastern margin of the Grand Banks, at the southern end of Flemish Pass. It is located 200 km southeast of the nearest cluster of earthquakes and has a distinctly different location from the other Grand Banks earthquakes in 1992. It lies under the continental slope, where the crustal thickness is about 10-15 km (see Discussion). Despite this, strong Lg-energy was observed on STJN.

#92d is a further earthquake in the cluster about 270 km east of St. John, the last of three in 1992.

"Note added in proof" The following four earthquakes occurred after the preparation of the figures, tables, and text of this report were completed in 1992. They are referenced only here and in the Appendix.

#94a is a M_L 3.2 earthquake on 940108 near the northeast trend of earthquakes discussed below under "Northeast Newfoundland Shelf and Basin".

#94b is a m_{bLg} 3.1 earthquake 35 km north of St. John's on 940811. It occurred during daylight hours, but despite numerous enquiries, no blasting source was identified. #69 was a similar-sized earthquake at a similar epicentral distance, but the unclear Sg arrival at STJN and the poor distribution of seismograph stations prevents a decision whether or not both occurred in the effectively same place. A subsequent event occurred on 950122 (#95).

#94c is a m_{bLg} 2.6 earthquake on 941201, about 80 km east of St. John's, distinctly farther east of St. John's than #69 or #94b.

#95 is a small, m_{bLg} 2.4, event on 950122, 35 km from St. John's, and likely in the same place as #69 and #94b. It occurred at night, so blasting can probably be ruled out.

DISCUSSION

Revision of epicentres

Both of the two earliest earthquakes originally placed on the Grand Banks have proved to be mislocated northern events. #22 gives a poorly constrained location in central Baffin Island, but likely occurred in the northern of the two seismically-active regions along the coast of Baffin Island (and not actually in the interior where the crude epicentre places it). In that region, its magnitude is not exceptional. #62 is confirmed as a Labrador Ridge event, and is one of several magnitude 5 events on the ridge. It would likely not have been mislocated if it had occurred a year later, when FBC had begun operating.

Of the other five earthquakes in the pre-1985 CEEF (Table 1; Fig. 2), one (#71) moves 32 km, two move about 17 km, and two less than 6 km (Table 4). Fifteen new events have been added to the CEEF since 1984 (11 given in Table 3, plus 4 in 94/95). Their CEEF epicentres reflect the incorporation of our enhanced research effort into the ongoing compilation of seismicity. Past experience suggests that without this enhanced effort some of these earthquakes would have been missed, and the routinely-determined epicentres for others would have needed revision.

Four earthquakes between 1975 and 1982 are reported for the first time, having been found from examination of the Memorial seismograms. During the same period only one earthquake (#76) had been located in the study area during the routine analysis. It is disturbing that three of the new earthquakes are larger than M_L 4 (4.0 to 4.2) and the closest (#81) is only 380 km from STJ (but 1280 km from the next closest seismograph of the Canadian Network, HAL). Although the amplitude on STJ was easily detectable, #81 gave less than 0.8 mm amplitude on all the other GSC seismograms.

Detection and location thresholds, and rates of activity

Now that the re-examination of the Memorial seismograms is complete it seems likely that since 1977 almost all earthquakes on the Grand Banks have been **detected** if they exceed the magnitude thresholds for the following radial distances east of St. John's:

| | | |
|--------|-----|-------|
| 350 km | 3.0 | m_N |
| 550 km | 3.5 | M_L |
| 900 km | 4.0 | M_L |

However, for a **location** requiring, say, HAL or SCH, the threshold has been considerably higher:

| | | |
|--------|-----|-------|
| 350 km | 4.0 | m_N |
| 550 km | 4.2 | M_L |
| 900 km | 4.6 | M_L |

Note that these thresholds are subjective estimates of a 90% confidence ability, based on the appearance of past earthquakes on the seismograms together with the attenuation rate implied by the M_L formula. A more rigorous analysis is intended in the context of a Canada-wide assessment of detection capability.

The location capability improved in 1983 with the opening of station GBN in Nova Scotia, and improved slightly for the period 1981-1991 because of the operation of four ECTN stations in New Brunswick. However, it is clear that smaller earthquakes on the inner Grand Banks, and larger earthquakes along the outer continental margin are not all being detected, and it is difficult to estimate how many are being missed, since the known annual rate of earthquakes in a particular magnitude range and in any given region is highly variable.

In addition, for the Grand Banks, it is difficult to calculate the annual probability of earthquakes larger than a given size because the location capability varies so dramatically in space and has varied also with time. If, however most of the magnitude 4 and greater earthquakes that occurred within 400 km of St. John's since 1977 are assumed to have been located, then Table 4 shows four earthquakes in this category between 1977 and 1994, i.e. four earthquakes in 18 years, or an average rate of 0.2 earthquakes of magnitude 4 and greater per annum on the inner Grand Banks. Since the located number is so small, then the annual rate is only a very rough estimate. If even two earthquakes in this category had been missed, then the annual rate would be 50% higher. Since there are even fewer earthquakes located with magnitudes greater than 4.5, they provide no direct basis for an annual rate of higher magnitude earthquakes east of St. John's. To make such estimates requires assumptions about the relative rates of big and small earthquakes; such assumptions are commonly made for input to seismic hazard calculations.

Lg Attenuation and propagation

There are numerous references in the brief earthquake descriptions above regarding the extent to which Lg-wave energy has propagated from the earthquake to St. John's. Lg-wave energy is known to be very sensitive to crustal structure, propagating very efficiently in normal continental crust but very poorly in oceanic or thinned-continental crust.

Figure 8 uses thick, moderate, and thin lines to show the paths with excellent, weakened, and zero Lg-wave propagation to St. John's. There are still too few paths to fully delimit the region of normal crust, though the 20-km contour of crustal thickness (see Fig. 10) appears to divide the earthquakes correctly. The difference in Lg propagation between #88b and that from the nearby earthquakes could be due to local crustal variations or to earthquake depth, but must be considered unexplained at this time.

Earthquakes with significant Lg-wave energy at St. John's commonly also show Lg-wave energy at Schefferville, but almost never in Nova Scotia or southern New Brunswick. The later path contains oceanic-thickness crust south of Newfoundland that attenuates the Lg-wave.

Focal Mechanisms (J. Adams has sole responsibility for the contents of this section)

None of the earthquakes discussed in this report is large enough to determine a reliable focal mechanism at present. This is mostly due to the extremely poor station distribution for these offshore earthquakes, as #71 would have been easily large enough for a reliable mechanism, had it occurred onshore.

At m_bL_g 4.8, #71 was large enough to be recorded by 84 seismographs reporting to the International Seismological Centre. Thirteen (3 being Canadian) reported polarities. Both Adams and Wahlström independently read all Canadian seismograms, confirmed the 3 Canadian readings reported to the ISC, and added 10 more (Table 4). Figure 9A shows the distribution of polarities. Of particular note is that STJ and HAL have opposite polarities, despite having similar azimuths from the epicentre. Figure 9B shows a range of mechanisms that fit the polarity data acceptably well (each misfits the equivalent of 4 full-weight polarities) without misfitting either STJ or HAL. These mechanisms, which represent almost pure strike-slip faulting on north-south or east-west planes, misfit our Canadian compressional readings at OTT, GWC, BLC, SES, and FFC and the ISC-reported polarity at ELT. Of these only the misfit of the clear polarity at FFC is of concern, but as the reported phase arrives 5 s late for GSC's solution, it may well be a depth phase rather than the direct P-wave arrival.

Two points to note about this poor mechanism are:

- pure strike-slip faulting is very uncommon in onshore southeastern Canada.
- the P-axis is oriented NW-SE, at a high angle to the "regional" stress field orientation which is NE-SW (Adams and Bell, 1991, Figure 3). However, local stress orientations estimated from oil-well breakouts in the Hibernia oilfield (70 km to the southeast) include anomalous east-west and north-south compression directions (Adams and Bell, 1991, Figure 16).

I suggest that future modelling (for example, of the Pnl wave on STJ and the teleseismic P-waves at Canadian Arctic stations) should be carried out to test my tentative mechanism and also to establish the depth of the 1971 earthquake.

Polarities from other earthquakes in this report are given in the PIK files (Appendix). Worthy of note is that STJ and MUNF are dilatational for Pn arrivals and HAL and GBN are

usually compressional, and that both compressional arrivals at STJ are for upward-directed arrivals from the close events #69 and #94b.

Seismotectonics

Figure 10 shows the relocated earthquakes together with bathymetry and crustal thickness contours taken from Shih et al. (1988). We discuss the earthquakes in five groups.

Mid-Atlantic Ridge. Event #89a lies in the extreme northeast of the study area, near the Mid-Atlantic Ridge. These earthquakes have no seismic hazard implications for Canada.

Newfoundland Ridge. Two earthquakes (#82 and #90) lie south of the Tail of the Bank, on the Newfoundland Ridge. Little is known of the seismotectonics of this region, though the pair of earthquakes might be associated with the rifted continental margin (Adams and Basham, 1989) and might represent reactivation of the extensional faults formed during the opening of the Atlantic Ocean.

Newfoundland Rifted Margin. A single earthquake (#92c) is known to have occurred under the continental slope that marks the southeastern edge of the Grand Banks. However, few earthquakes smaller than M4 can be located in this area. It occurred in a region of continental crust somewhat thinned by the rifting that formed the Atlantic Ocean. Although this is the first known earthquake from the 1000-km-long rifted margin that bounds the Grand Banks to the southeast and southwest, seismicity of the margin was predicted by the 'ESX' seismicity model of Basham and Adams (1983) and Basham et al. (1983). Their model suggested that the entire rifted margin would be capable of generating large ($M \approx 7$) earthquakes. By implication, either the entire margin would need to be seismically-active at a low level (below the detection threshold) or earthquakes were occurring at long intervals, so that contemporary seismicity (which at the time was absent) might not reveal all active structures. Therefore #92c is important supporting evidence for the ESX model, which itself is important because of the seismic hazard large earthquakes on the margin would pose to facilities producing hydrocarbons from the Jeanne D'Arc Basin.

Grand Banks Proper. Seven earthquakes (#71, #88b, #89b, #91, #92a, #92b, and #92d) lie on the Grand Banks in the vicinity of the Hibernia Oil Field. #91 lies about 70 km to the west of the other six, but the azimuthal precision is such that the remaining six earthquakes (except, perhaps #89b) could well have occurred in the same place (Fig. 7). Although not conclusive, an analysis of relative arrival times for Nova Scotia and Quebec stations (representing the maximum azimuthal spread of the data) suggests a NNE-SSW to N-S alignment of the epicentres. This possible alignment roughly parallels the structural trends on the central Grand Banks.

The epicentres lie on the Bonavista Platform, about 20 km west of, but parallel to, the Mercury Fault (Adams and Bell, 1991, Figure 16). In view of the potential location bias (due to the unknown crustal velocity structure), it is not impossible that they actually occurred on the Mercury Fault. Alternatively, they may have occurred on a less-significant or an unknown fault within the Bonavista Platform. The Mercury Fault is an east-dipping listric detachment which bounds the northwest side of the Jeanne D'Arc Basin, the largest of a number of northeast-trending basins that cut across the Grand Banks. These basins were formed by normal faulting of the basement consequent on crustal stretching during the opening of the Atlantic Ocean (Ziegler, 1988).

This cluster of activity represents the most active region on the Grand Banks, having generated three M4, and four M3 earthquakes. Allowing for different periods of completeness, the recorded earthquakes suggest an annual rate for $M \geq 4.0$ earthquakes of about 0.1, which makes the cluster about a third as active as the Laurentian Slope zone south of Newfoundland (Keen et al., 1990; Adams, unpub.).

East Newfoundland Basin and Shelf. One earthquake (#81) lies south of Orphan Knoll and 200 km northeast of the cluster of earthquakes near Hibernia. The remaining nine earthquakes lie along a northeast trend extending from just off St. John's (#69) to 52°N, 47°W. The earthquakes in this trend lie in a band about 70-100 km wide, and might extend southwest to include earthquakes off Bonavista in 1965 and on the northern Avalon Peninsula in 1884 and 1956 (Adams and Staveley, 1985).

The earthquakes near the middle of the trend (#76, #77, #84a,b, #85, #88a) lie in a region where the basement surface has a local relief of 2-4 km and is formed into linear, northeast-trending ridges (Grant, 1988). These are interpreted as the result of

strong block faulting of the Paleozoic basement in Jurassic times (Proctor et al., 1984, p. 39).

The two most northeasterly earthquakes (#70 and #78) occurred on oceanic crust close to the inshore extension of the Charlie-Gibbs Fracture Zone. Smaller earthquakes could seldom be located in this area. In this region the fracture zone trends east-northeast. In the Labrador Sea to the north, Adams and Simmons (1991) have found that many of the earthquakes on the oceanic crust occur in the vicinity of the major fracture zones.

In summary, the seismicity discussed in this report may be due to fracture zones weakening the oceanic crust around the Grand Banks, or to northeast-trending normal faults and rift features that broke the integrity of the basement in the Mesozoic during the opening of the Atlantic Ocean. At this point in our understanding there are no correlations with individual faults, only spatial associations with broad zones of weakness. These conclusions are not very different from those of Adams and Basham (1989).

Seismic Hazard Implications

Earthquakes on the Grand Banks have significant seismic hazard implications for offshore hydrocarbon production platforms such as will be emplaced at Hibernia or elsewhere on the Grand Banks. Some earlier hazard estimates included the two largest earthquakes (1922 and 1962) in the hazard calculations, but these earthquakes have now been relocated away from the Grand Banks. The largest earthquake is now m_bLg 4.8, but more importantly it is now seen to lie within an active cluster, revealed chiefly because of activity in the past five years.

The problems with detection and location thresholds (discussed above) make it difficult to assess the rates of earthquakes and hence the hazard for the offshore. Previous hazard estimates (Basham et al., 1983; Adams, 1986) recognized the poorly determined nature of the offshore seismicity and the need to relocate the older earthquakes and to improve seismograph coverage to detect more of the small offshore earthquakes. Some relatively inexpensive additions to the onland seismograph network – for example a station near Bonavista, Cape Freels, or St. Anthony – would certainly improve the detection and location of offshore

earthquakes. Although expensive, and technically difficult to install and operate, a tethered ocean bottom seismometer in the vicinity of the Hibernia platform would present the best near-term opportunity to improve seismic monitoring on the outer Grand Banks. One further conclusion of the present report is that effort needs to be made to assess the earthquake location threshold in both time and space so that the seismicity record can be interpreted in terms of recurrence rates for the entire Grand Banks.

CONCLUSIONS

1. The two largest earthquakes once considered on the Grand Banks have been relocated out of the area. The largest earthquake is now m_{bLg} 4.8.
2. Four new earthquakes between 1977 and 1985 have been found, the epicentres of five previously-known earthquakes prior to 1985 have been revised slightly, and post-1984 records have been searched for earthquakes with more care, resulting in a further 15 earthquakes being added.
3. Most earthquakes appear to be spatially associated with Mesozoic extensional structures that trend northeastwards across the Grand Banks, although another decade of monitoring may be needed to test this conclusion.
4. A cluster of activity northwest of the Hibernia site represents the most active region on the Grand Banks in recent years.
5. The earthquakes catalogued in this report form the most reliable basis for seismic hazard calculations for the Grand Banks.

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TABLE CAPTIONS

Table 1. 1922 to 1984 earthquake epicentres, as known in 1985.

Table 2. Probable earthquakes recorded on Memorial seismographs, but unlocatable. (possible Grand Banks earthquakes)

Table 3. A. New solutions for earthquakes no longer in the study area. B. Revised, new, & added solutions for earthquakes in the study area.

Table 4. Summary of revisions to Grand Banks earthquakes.

Table 5. Polarity data for the 1971 Grand Banks earthquake.

FIGURE CAPTIONS

Figure 1. Map of the Eastern Margin showing the geographic scope of the two published reports and the region of the present study.

Figure 2. Seismicity of the Grand Banks as known in 1985, taken from the then-current Canadian Earthquake Epicentre File. Only earthquakes within the study area (east of the heavy line) are shown. Symbols in this and subsequent figures are:— star, magnitude (M) ≥ 5.0 ; square, $4.0 \leq M \leq 4.9$; triangle, $3.0 \leq M \leq 3.9$; and cross, $M \leq 3.0$.

Figure 3. Relocated Grand Banks earthquakes. The label next to each earthquake symbol identifies the earthquake by year (76, 81, etc) and sequence in the year (84a, 84b, etc) and reference to Table 3.

Figure 4. Relocation vectors for the studied earthquakes. The heavy lines extend back from the revised epicentre (also marked by a symbol) to the pre-revision epicentre. For most earthquakes the revisions are small, less than the symbol size, but two events have moved north, off the map area (see text).

Figure 5. Computed uncertainty for the revised epicentres, as represented by precision diamond (the axes represent the standard errors in latitude and longitude). For almost all the earthquakes, the computed precision is less than the space occupied on the map by the earthquake symbol (see also Fig. 6 and 7), though the plotted epicentres may be inaccurate by an amount at least as large as the symbol size.

Figure 6. Detail of Figure 5, showing the seismicity of the Grand Banks proper, together with the precision diamonds.

Figure 7. Detail of Figure 5, showing the epicentres in the cluster of earthquakes near Hibernia, together with their precision diamonds.

Figure 8. Propagation of Lg-wave energy to St. John's from offshore earthquakes. Thick lines represent paths with efficient propagation, single moderate-thickness line represents partial attenuation of Lg, and thin lines represent paths that completely attenuate Lg.

Figure 9. Focal mechanism information for event #71. A) Data (see Table 5). C and D represent full-weight data, c and d represent half-weight data. B) Possible nodal planes, representing strike-slip faulting on a N-S or an E-W plane. P, B, and T represent the axes of the solution.

Figure 10. Relocated offshore earthquakes, crustal thickness (taken from Shih et al., 1988), and 500-m isobath. Compare with Figure 3 to determine the identity of the events. As in previous maps, earthquakes west of the heavy line are not plotted.

TABLE 1

1922 TO 1984 EARTHQUAKE EPICENTRES AS KNOWN IN 1985

| REF | DATE | TIME (UT) | LAT | LONG | DEPTH (km) | MAGNITUDE | TYPE |
|-----|----------|--------------|---------|---------|---------------|-----------|------|
| ISC | 19220726 | 062855 | 50.0 N | 50.0 W | | 5.3 | ML |
| GSC | 19620803 | 013102 | 52. N | 54.2 W | | 4.8 | ML |
| GSC | 19690805 | 215324 | 47.74 N | 52.31 W | 18 | 3.3 | MN |
| GSC | 19701031 | 174445 | 52.17 N | 46.37 W | 18 | 4.3 | ML |
| ISC | 19710815 | 061715 | 47.46 N | 49.53 W | | 4.8 | MB |
| GSC | 19760828 | 192330 | 50.10 N | 48.85 W | 18 | 4.0 | ML |
| GSC | 19840526 | 192549 | 48.89 N | 51.06 W | 18 | 3.5 | MN |

REF: GSC - solution from GSC, ISC - solution from ISC

TABLE 2

PROBABLE EARTHQUAKES RECORDED ON MEMORIAL
SEISMOGRAPHS, BUT UNLOCATABLE.
(POSSIBLE GRAND BANKS EARTHQUAKES)

| DATE | TIME UT | STATION | S-P (s) | DIST (km) | AMP (mm) | MAGNITUDE ¹ ML |
|--------|------------|---------|------------|--------------|-------------|------------------------------|
| 750403 | 0941 | MUNF | 47 | 470 | 2.0 | 3.4 |
| 760217 | 0658 | MUNF | 45 | 450 | 3.0 | 2.9 |
| 760327 | 1101 | MUNF | 63 | 645 | 2.5 | 3.8 |
| 761124 | 2154 | MUNF | 160 | 1710 | 7.0 | - |
| 770811 | 1048 | MUNF | 42 | 420 | 2.5 | 3.5 |
| 780618 | 1316 | CBK | 69 | 710 | - | - |
| 780228 | 0756 | MUNF | 75 | 780 | 2.0 | 3.7 |
| 790408 | 2132 | MUNF | 43 | 425 | 1.5 | 3.4 |
| 790712 | 0314 | MUNF | 67 | 690 | 1.5 | 3.6 |
| 790725 | 0444 | MUNF | 62 | 635 | | |
| 791014 | 1407 | MUNF | 24 | 220 | 4.0 | 2.9 |
| 791112 | 1429 | MUNF | 32 | 305 | 6.0 | 3.0 |
| 800731 | 1609 | MUNF | 82 | 850 | - | - |
| 810121 | 0426 | CBK | 61 | 625 | - | - |
| 821215 | 0654 | MUNF | 23 | 205 | - | - |
| 920218 | 0346 | STJN | 54 | 550 | 0.8 | 3.0 |
| 920925 | 1818 | STJN | 63 | 645 | 0.3 | 2.9 |

¹Magnitude assuming Sn amplitudes; '-' not determined

TABLE 3

A. NEW SOLUTIONS FOR EARTHQUAKES NO LONGER IN THE STUDY AREA.

| ID ¹ | ST | LAT | LONG | MAG | TIME | DATE | ERRORS | STN | PH | M | RMS | DEPTH |
|-----------------|----|------------------|--------|--------|---------|----------|-------------|-----|----|---|-----|-------------|
| 22 | M | +70.000- | 75.000 | MS=5.1 | 0628550 | 26071922 | 00.0000.000 | 0.0 | 2 | 3 | 0 | 0.00N 18.00 |
| | | (Baffin Island) | | | | | | | | | | |
| 62 | M | +61.027- | 58.225 | ML=4.8 | 0131042 | 03081962 | 00.0000.000 | 0.0 | 2 | 3 | 1 | 0.00 18.00 |
| | | (Labrador Ridge) | | | | | | | | | | |

B. REVISED, NEW, & ADDED SOLUTIONS FOR EARTHQUAKES IN THE STUDY AREA.

| ID ¹ | ST | LAT | LONG | MAG | TIME | DATE | ERRORS | STN | PH | M | RMS | DEPTH |
|-----------------|----|----------|--------|--------|---------|----------|-------------|-----|----|----|-----|-------------|
| 69 | R | +47.632- | 52.156 | MN=3.4 | 2153207 | 05081969 | 00.0860.110 | 0.2 | 5 | 9 | 4 | 0.94 18.00 |
| 70 | R | +52.189- | 46.333 | ML=4.3 | 1744440 | 31101970 | 00.0310.110 | 0.3 | 2 | 4 | 2 | 0.29 18.00 |
| 71 | R | +47.308- | 49.160 | MN=4.8 | 0617095 | 15081971 | 00.1310.215 | 0.2 | 10 | 17 | 6 | 2.50 18.00 |
| 76 | R | +50.133- | 48.910 | ML=4.1 | 1923304 | 28081976 | 00.0260.108 | 0.3 | 5 | 8 | 3 | 0.54 18.00 |
| 77 | N | +49.337- | 49.586 | ML=3.4 | 0805261 | 25031977 | 00.1230.155 | 0.0 | 2 | 4 | 1 | 0.12 18.00 |
| 78 | N | +52.573- | 46.752 | ML=4.1 | 1435140 | 03011978 | 00.0280.073 | 0.1 | 3 | 6 | 3 | 0.36 18.00 |
| 81 | N | +48.643- | 47.820 | ML=4.2 | 1306126 | 26091981 | 00.0670.170 | 0.2 | 5 | 10 | 6 | 1.16 18.00 |
| 82 | N | +41.367- | 47.282 | ML=4.0 | 0407239 | 22071982 | 00.1040.124 | 0.1 | 6 | 8 | 4 | 1.05 18.00 |
| 84a | R | +48.852- | 50.837 | MN=3.6 | 1925475 | 26051984 | 00.0190.037 | 0.3 | 8 | 18 | 6 | 0.48 18.00 |
| 84b | A | +48.852- | 50.837 | MN=2.6 | 2204198 | 26051984 | 00.0000.000 | 0.0 | 1 | 1 | 0 | 0.00H 18.00 |
| 85 | A | +49.787- | 50.480 | ML=2.8 | 1806035 | 03081985 | 00.0150.053 | 0.3 | 4 | 6 | 2 | 0.14 18.00 |
| 88a | A | +49.834- | 49.982 | ML=3.5 | 0715448 | 09011988 | 00.0180.058 | 0.2 | 8 | 11 | 6 | 0.34 18.00 |
| 88b | A | +47.463- | 49.246 | ML=3.1 | 0016451 | 09081988 | 00.1950.120 | 0.0 | 5 | 8 | 2 | 0.76 18.00 |
| 89a | A | +52.730- | 35.200 | MS=5.4 | 2253371 | 14051989 | 00.0390.030 | 0.0 | - | - | 25 | 0.00H 10.00 |
| 89b | A | +46.960- | 49.247 | ML=4.2 | 0515398 | 03121989 | 00.0500.088 | 0.4 | 6 | 9 | 4 | 0.71 18.00 |
| 90 | A | +41.918- | 48.341 | ML=4.7 | 2140041 | 24041990 | 00.0270.062 | 0.1 | 14 | 23 | 13 | 0.67 18.00 |
| 91 | A | +47.364- | 50.167 | MN=3.2 | 1123013 | 23071991 | 00.0820.146 | 0.2 | 6 | 10 | 3 | 0.79 18.00 |
| 92a | A | +47.245- | 49.236 | MN=4.0 | 0607283 | 13011992 | 00.0570.100 | 0.2 | 17 | 25 | 14 | 0.93 18.00 |
| 92b | A | +47.326- | 49.346 | ML=3.0 | 1658161 | 06071992 | 00.3710.268 | 0.2 | 3 | 6 | 3 | 0.74 18.00 |
| 92c | A | +46.118- | 47.438 | MN=3.9 | 0420227 | 17071992 | 00.1120.087 | 0.1 | 9 | 17 | 2 | 0.77 18.00 |
| 92d | A | +47.353- | 49.114 | MN=3.4 | 1131520 | 10081992 | 00.0830.070 | 0.0 | 5 | 9 | 1 | 0.39 18.00 |

¹This table is modified from GSC's CEEF format. A brief description is given by the column heads as follows: ID, see text; ST, status - M moved, R revised, N new, A added; LAT, LONG latitude and longitude in decimal degrees; MAG, magnitude; TIME, hour minute seconds in Universal Time; DATE; ERRORS, standard errors on latitude, longitude, and magnitude; STN, number of stations used for location; PH, number of phases used; M, number of stations used for average magnitude; RMS, root mean square residual (seconds); DEPTH, assigned depth (km).

TABLE 4

SUMMARY OF REVISIONS TO GRAND BANKS EARTHQUAKES

| ID | SEQ | DATE | TIME | Δ LAT | Δ LONG | Δ MAG | Δ Z | STATUS | Δ KM |
|-----|-----|---------------|------|--------------|---------------|--------------|------------|---------|-------------|
| 22 | 1 | 19220726.0628 | | 20.000 | -25.000 | -0.2 | 18.00 | moved | 2852.8 |
| 62 | 2 | 19620803.0131 | | 9.027 | -4.025 | 0.0 | 18.00 | moved | 1040.9 |
| 69 | 3 | 19690805.2153 | | -0.108 | 0.154 | 0.1 | 0.00 | revised | 16.6 |
| 70 | 4 | 19701031.1744 | | 0.019 | 0.037 | 0.0 | 0.00 | revised | 3.3 |
| 71 | 5 | 19710815.0617 | | -0.152 | 0.370 | -0.1 | 18.00 | revised | 32.5 |
| 76 | 6 | 19760828.1923 | | 0.033 | -0.060 | 0.1 | 0.00 | revised | 5.6 |
| 77 | 7 | 19770325.0805 | | 0.000 | 0.000 | 3.4 | 0.00 | new | 0.0 |
| 78 | 8 | 19780103.1435 | | 0.000 | 0.000 | 4.1 | 0.00 | new | 0.0 |
| 81 | 9 | 19810926.1306 | | 0.000 | 0.000 | 4.2 | 0.00 | new | 0.0 |
| 82 | 10 | 19820722.0407 | | 0.000 | 0.000 | 4.0 | 0.00 | new | 0.0 |
| 84a | 11 | 19840526.1925 | | -0.038 | 0.223 | 0.1 | 0.00 | revised | 16.8 |
| 84b | 12 | 19840526.2204 | | 0.000 | 0.000 | 2.6 | 0.00 | added | 0.0 |
| 85 | 13 | 19850803.1806 | | 0.000 | 0.000 | 2.8 | 0.00 | added | 0.0 |
| 88a | 14 | 19880109.0715 | | 0.000 | 0.000 | 3.5 | 0.00 | added | 0.0 |
| 88b | 15 | 19880809.0016 | | 0.000 | 0.000 | 3.1 | 0.00 | added | 0.0 |
| 89a | 16 | 19890514.2353 | | 0.000 | 0.000 | 5.4 | 0.00 | added | 0.0 |
| 89b | 17 | 19891203.0515 | | 0.000 | 0.000 | 4.2 | 0.00 | added | 0.0 |
| 90 | 18 | 19900424.2140 | | 0.000 | 0.000 | 4.7 | 0.00 | added | 0.0 |
| 91 | 19 | 19910723.1123 | | 0.000 | 0.000 | 3.2 | 0.00 | added | 0.0 |
| 92a | 20 | 19920113.0607 | | 0.000 | 0.000 | 4.0 | 0.00 | added | 0.0 |
| 92b | 21 | 19920706.1658 | | 0.000 | 0.000 | 3.0 | 0.00 | added | 0.0 |
| 92c | 22 | 19920717.0420 | | 0.000 | 0.000 | 3.9 | 0.00 | added | 0.0 |
| 92d | 23 | 19920810.1131 | | 0.000 | 0.000 | 3.4 | 0.00 | added | 0.0 |

ID and SEQ represent the earthquake identification number and a simple counting sequence, respectively; Δ LAT, Δ LONG, Δ MAG and Δ Z represent the changes to these parameters established in this Open File; STATUS describes the reason for the change: moved - moved out of the area; revised - existing solution improved; new - event found on old MUN seismograms; added - event added from enhanced reading of MUN and GSC seismograms; Δ KM represents the distance moved in kilometres.

TABLE 5

POLARITY DATA FOR THE 1971 GRAND BANKS EARTHQUAKE

| ID | AZ | TO | POL | SOURCE |
|-----|---------|--------|-----|--|
| STJ | 274.000 | 49.000 | D | ISC and Wahlstrom |
| SCH | 311.000 | 49.000 | - | Adams -, ISC D, Wahlstrom C |
| MNT | 272.000 | 46.000 | - | ISC and Adams |
| CLE | 267.000 | 33.000 | D | ISC |
| FAV | 267.000 | 29.000 | C | ISC |
| DOU | 65.000 | 29.000 | C | ISC |
| NOR | 8.000 | 28.000 | C | ISC |
| GOL | 280.000 | 28.000 | D | ISC |
| ILT | 340.000 | 23.000 | D | ISC |
| BNG | 103.000 | 19.600 | D | ISC |
| ELT | 26.000 | 19.200 | D | ISC |
| AAB | 37.000 | 18.000 | C | ISC |
| GAR | 44.000 | 18.000 | C | ISC |
| HAL | 261.000 | 49.000 | C | Wahlstrom and Adams good quality, not strong |
| SIC | 290.000 | 49.000 | - | Wahlstrom C, Adams - |
| OTT | 274.000 | 42.000 | + | very weak, noisy record |
| UNB | 270.000 | 49.000 | - | clear, not strong, Wahlstrom D |
| SFA | 277.000 | 49.000 | D | good |
| GWC | 305.000 | 39.000 | + | v weak on noisy record |
| SUD | 280.000 | 35.000 | E | low signal strength, emergent |
| BLC | 321.000 | 33.000 | + | Adams 900925 |
| FFC | 303.000 | 32.000 | C | Adams 900925 good polarity but 5 sec late |
| SES | 298.000 | 30.000 | + | Adams 900925 |
| INK | 329.000 | 29.000 | D | Adams 900925 |
| MCE | 303.000 | 29.000 | - | ISC |

ID, Station providing the reading; AZ azimuth from earthquake to station; TO take-off angle from earthquake to station; POL P-wave first motion:- C and D are compression and dilatation; + and - are half-weight C and D.

Eastern Margin Study Areas

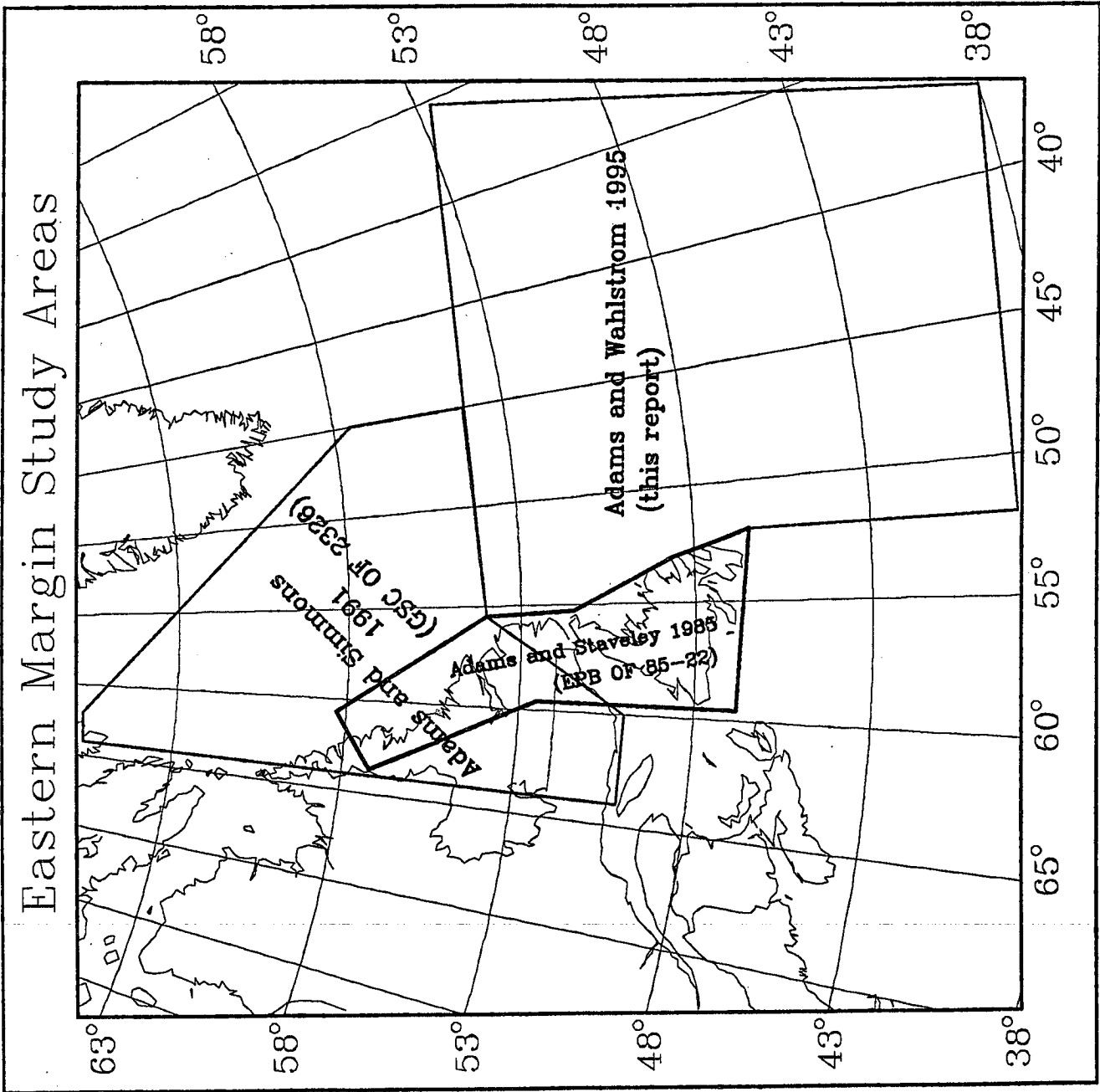


Figure 1. Map of the Eastern Margin showing the geographic scope of the two published reports and the region of the present study.

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DIVISION DE LA GEOPHYSIQUE COMMISSION GEOLOGIQUE DU CANADA

Figure 2. Seismicity of the Grand Banks as known in 1985, taken from the then-current Canadian Earthquake Epicentre File. Only earthquakes within the study area (east of the heavy line) are shown. Symbols in this and subsequent figures are:- star, magnitude (M) ≥ 5.0 ; square, $4.0 \leq M \leq 4.9$; triangle, $3.0 \leq M \leq 3.9$; and cross, $M \leq 3.0$.

DEFINITIONS

- x M ≤ 3
- ▲ M ≥ 3
- M ≥ 4
- ★ M ≥ 5

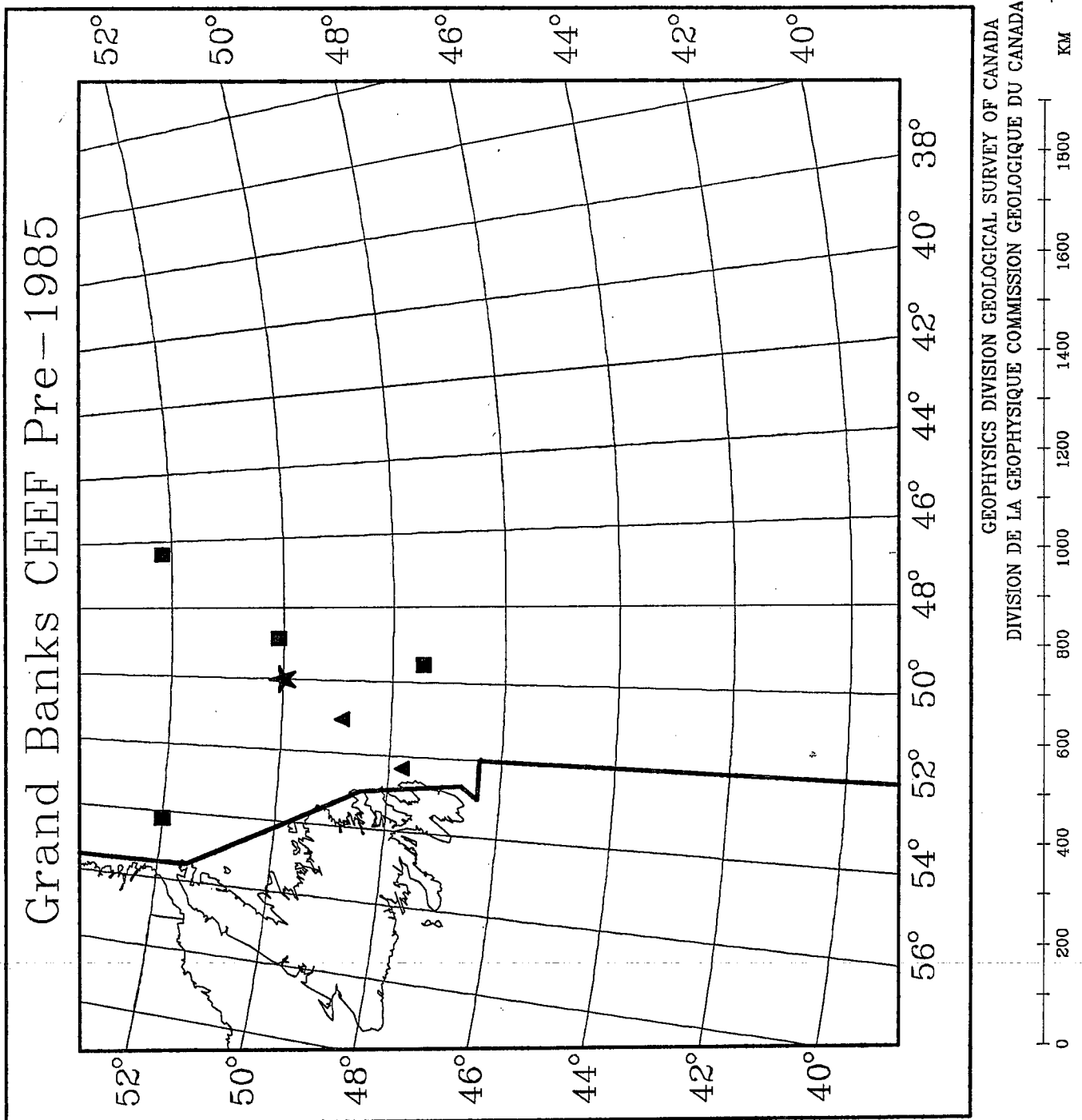


Figure 3. Relocated Grand Banks earthquakes. The label next to each earthquake symbol identifies the earthquake by year (76, 81, etc) and sequence in the year (84a, 84b, etc) and reference to Table 3.

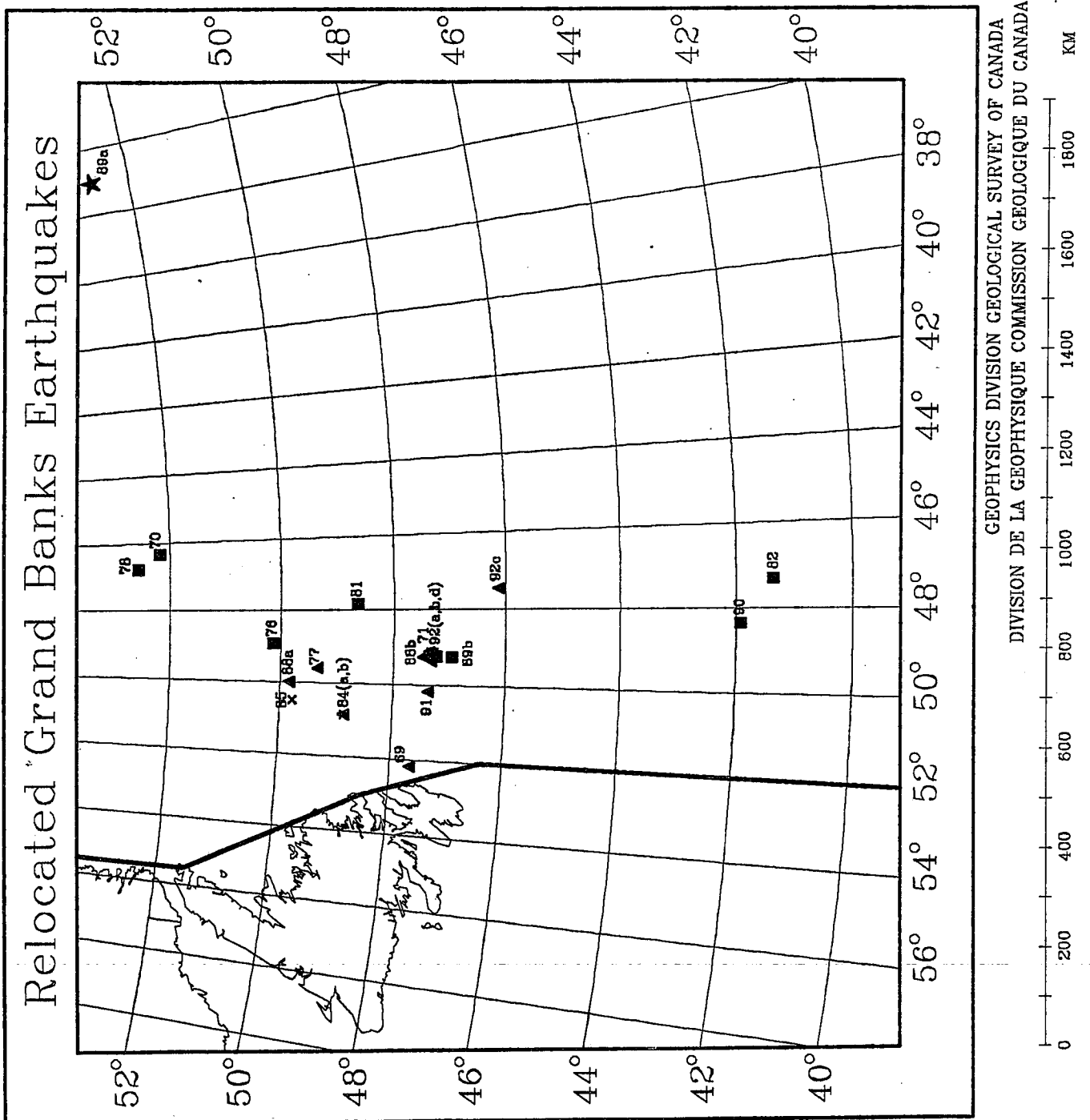


Figure 4. Relocation vectors for the studied earthquakes. The heavy lines extend back from the revised epicentre (also marked by a symbol) to the pre-revision epicentre. For most earthquakes the revisions are small, less than the symbol size, but two events have moved north, off the map area (see text).

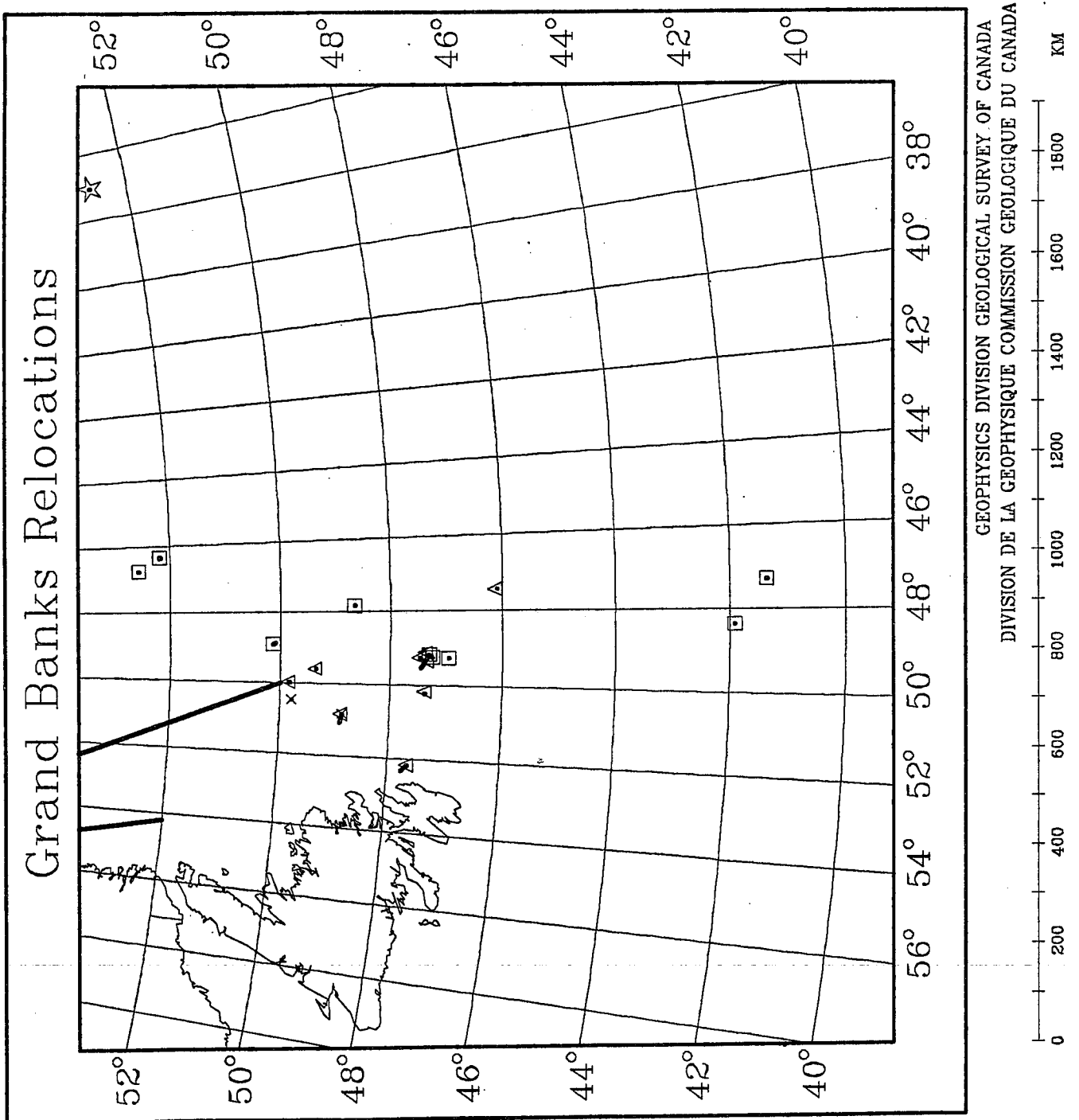
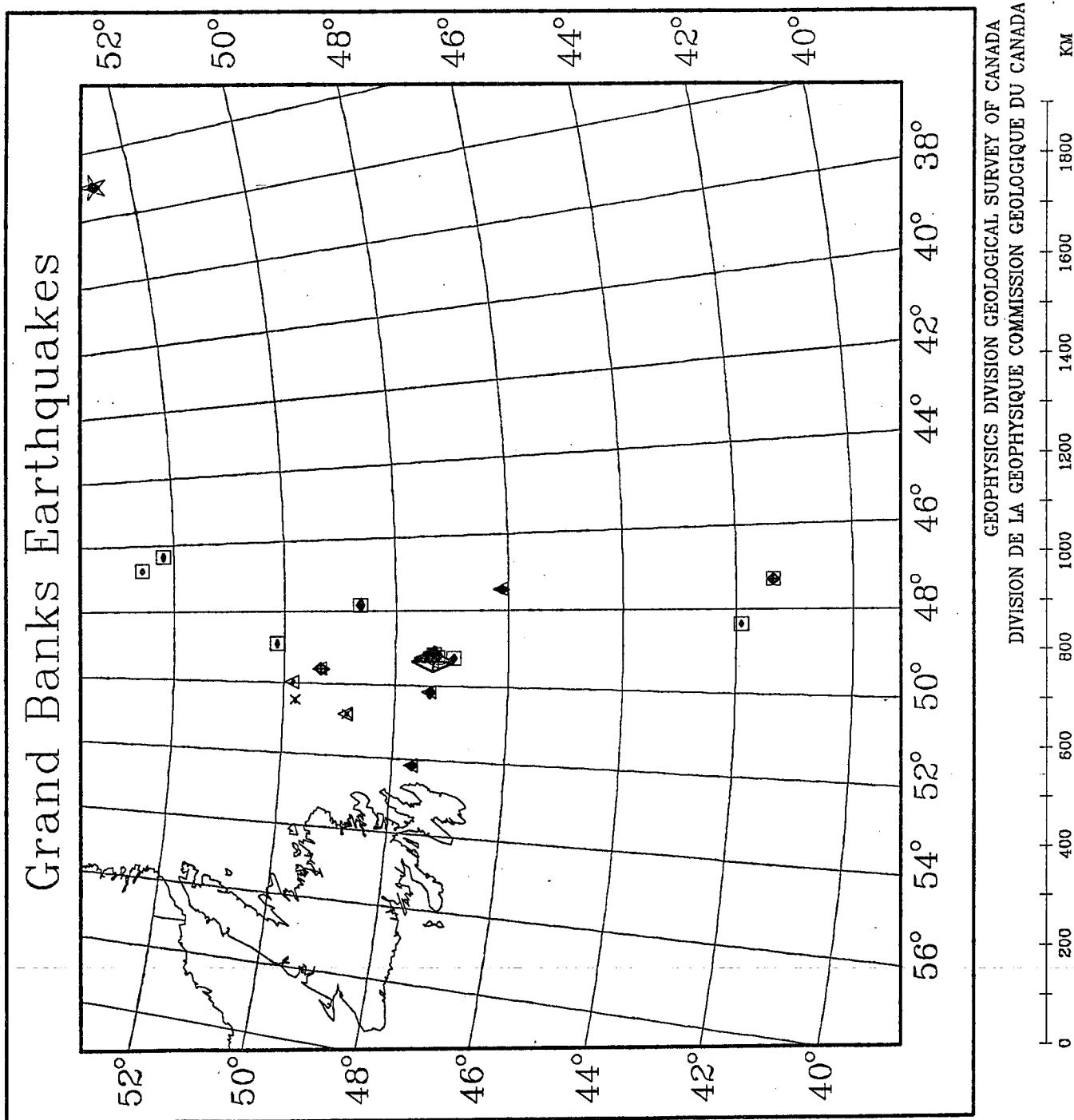


Figure 5. Computed uncertainty for the revised epicentres, as represented by precision diamond (the axes represent the standard errors in latitude and longitude). For almost all the earthquakes, the computed precision is less than the space occupied on the map by the earthquake symbol (see also Fig. 6 and 7), though the plotted epicentres may be inaccurate by an amount at least as large as the symbol size.



010406
48.75N
51.69W
3.1MN

Figure 6. Detail of Figure 5, showing the seismicity of the Grand Banks proper, together with the precision diamonds.

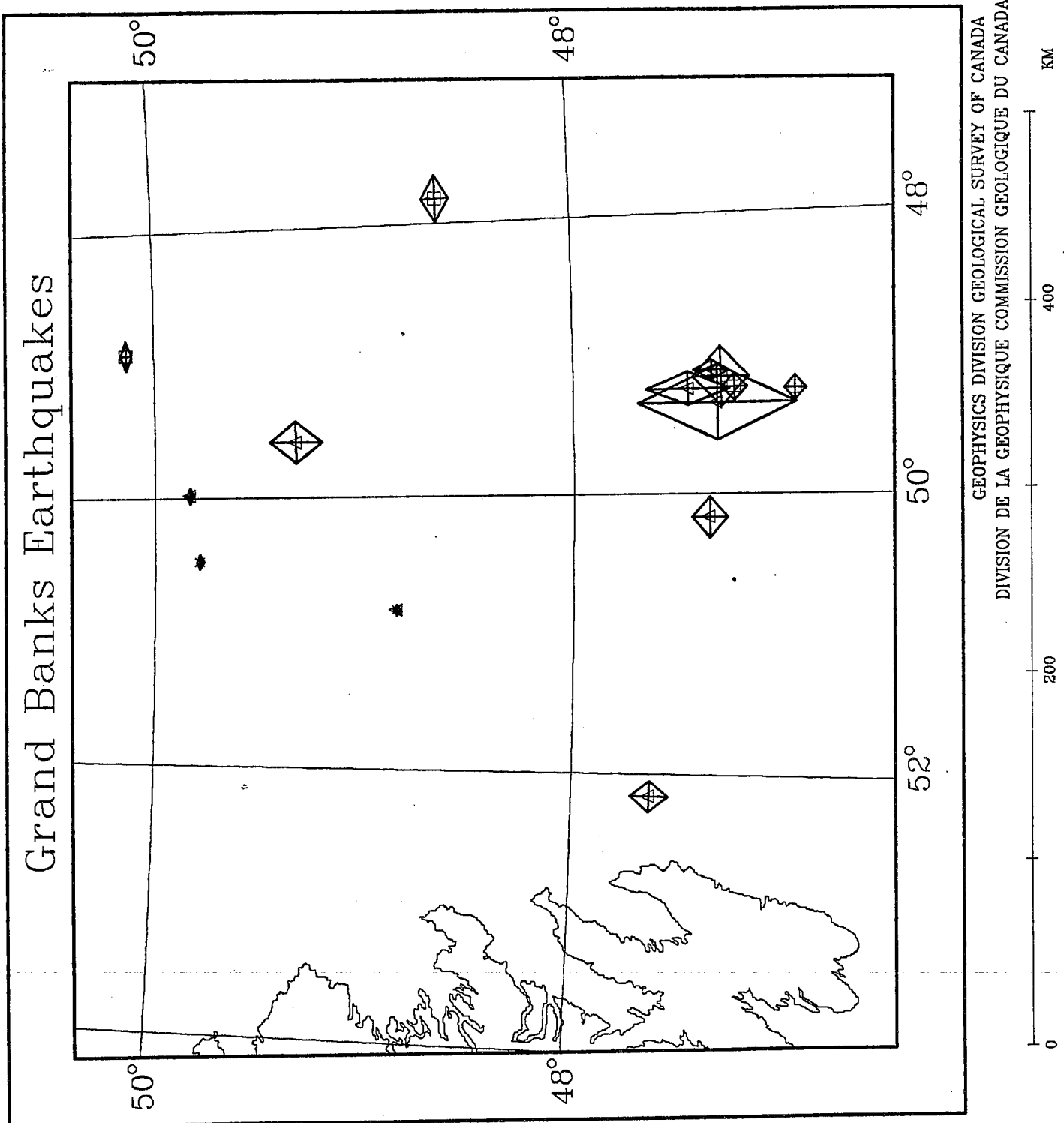


Figure 7. Detail of Figure 5, showing the epicentres in the cluster of earthquakes near Hibernia, together with their precision diamonds.

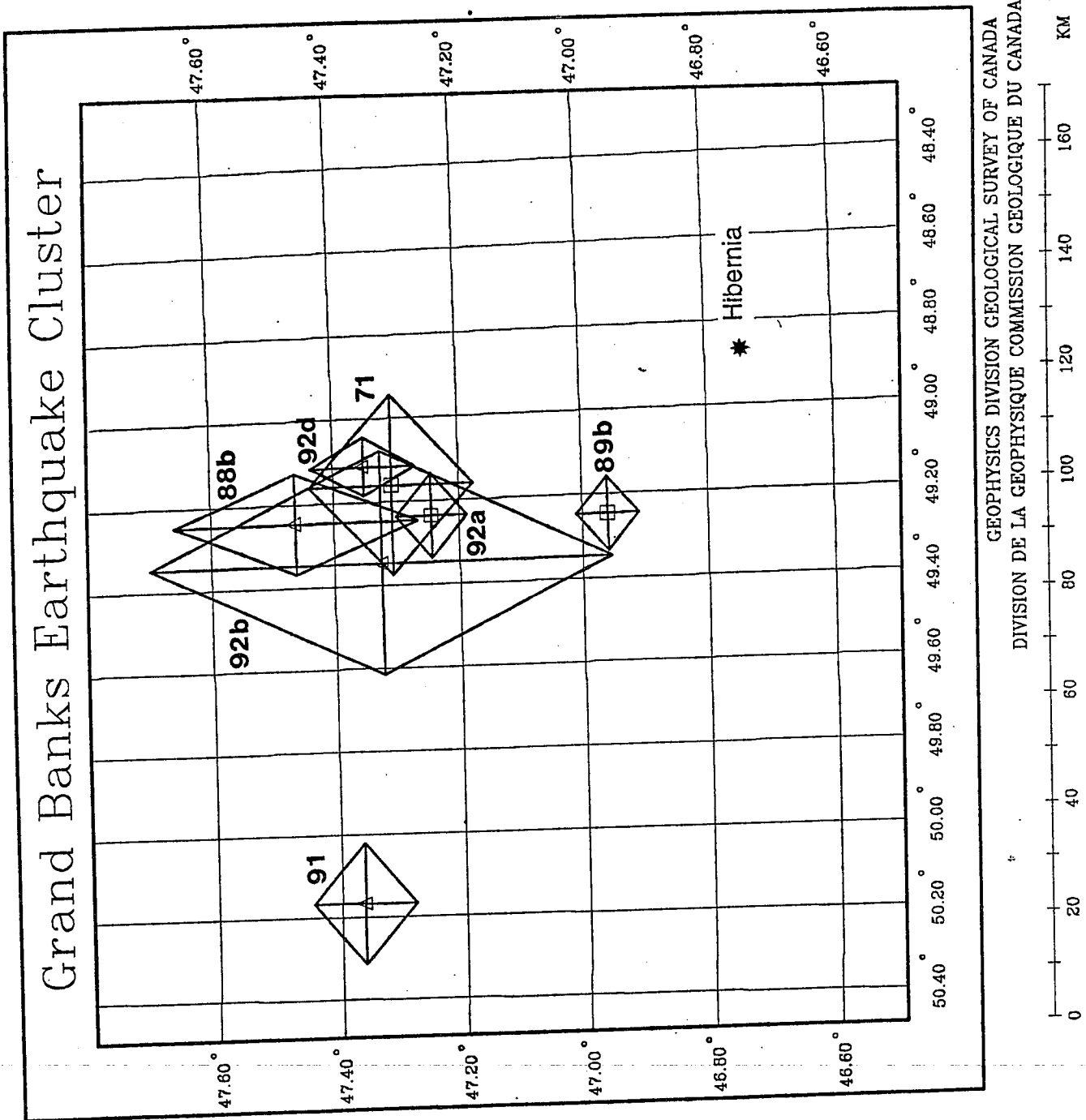
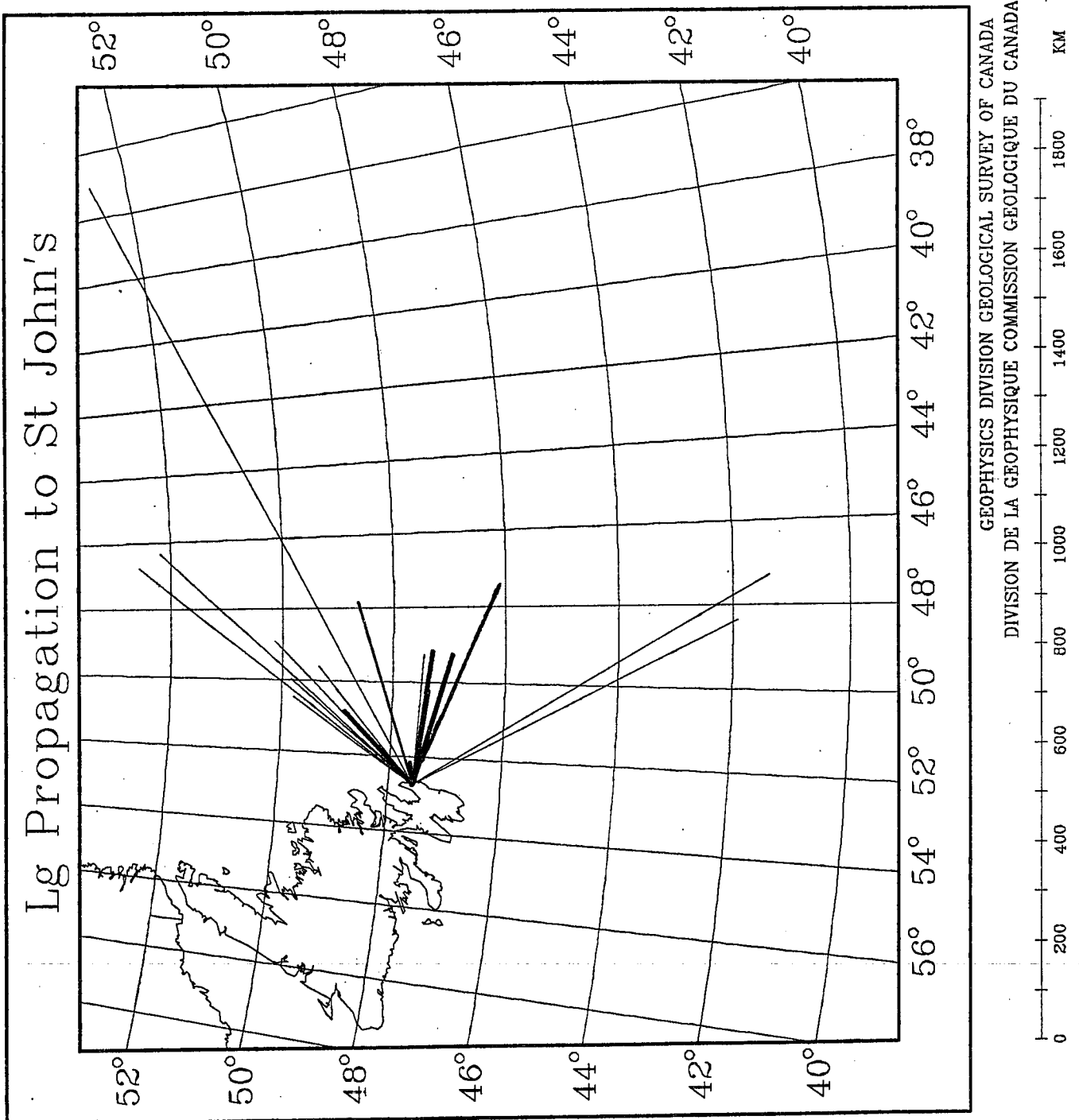


Figure 8. Propagation of Lg-wave energy to St. John's from offshore earthquakes. Thick lines represent paths with efficient propagation, single moderate-thickness line represents partial attenuation of Lg, and thin lines represent paths that completely attenuate Lg.



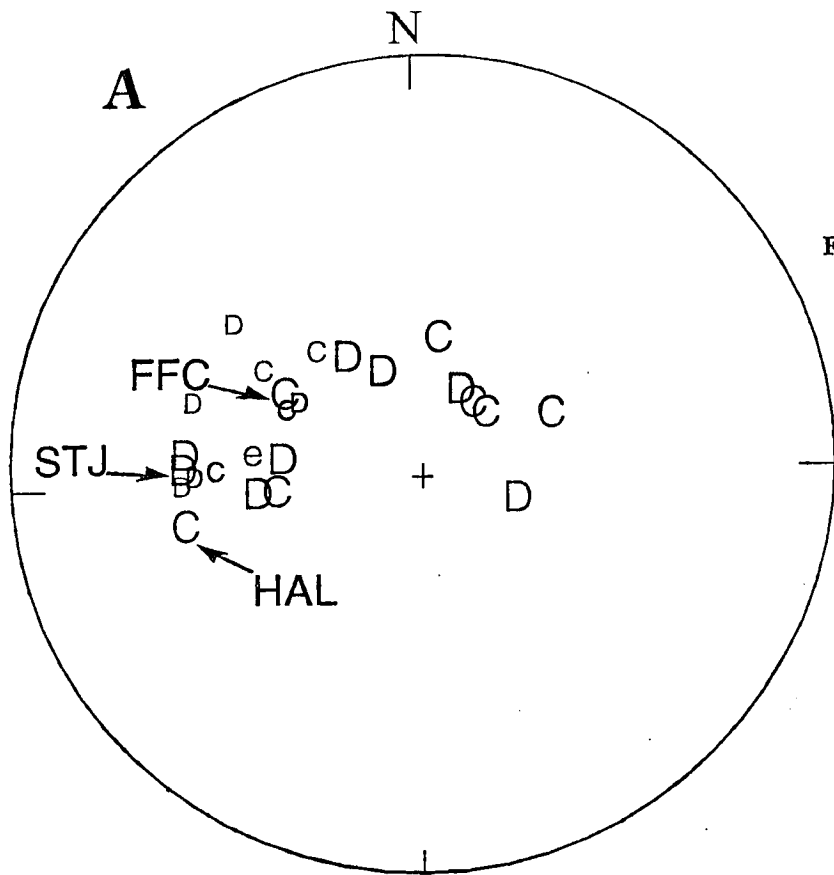


Figure 9. Focal mechanism information for event #71. A) Data (see Table 5). C and D represent full-weight data, c and d represent half-weight data. B) Possible nodal planes, representing strike-slip faulting on a N-S or an E-W plane. P, B, and T represent the axes of the solution.

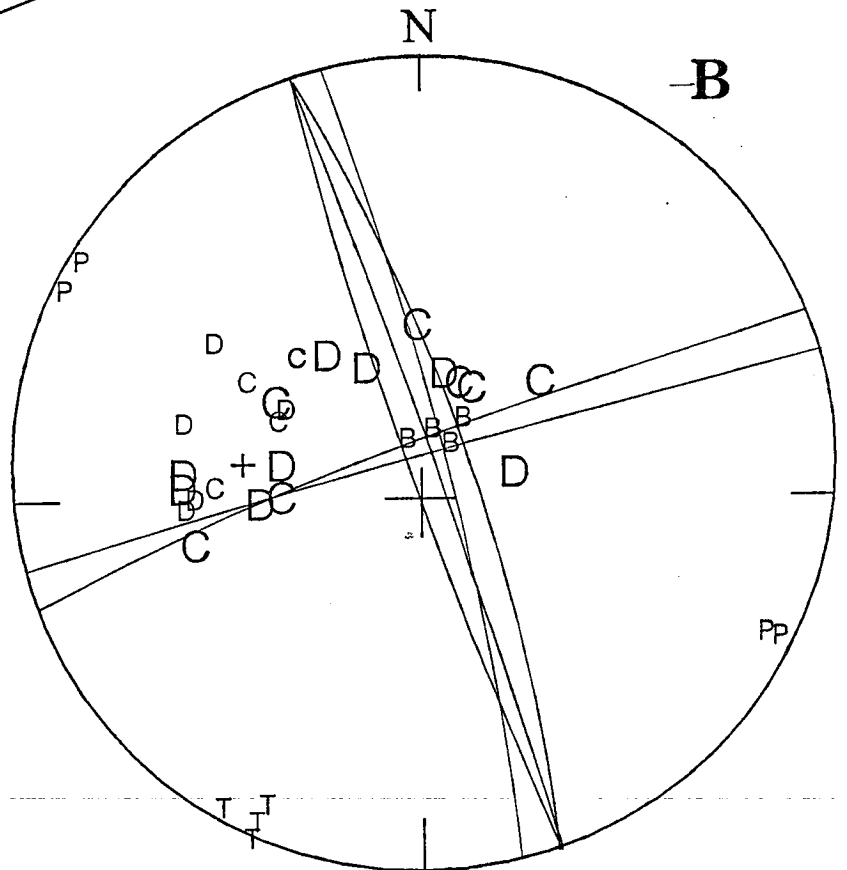
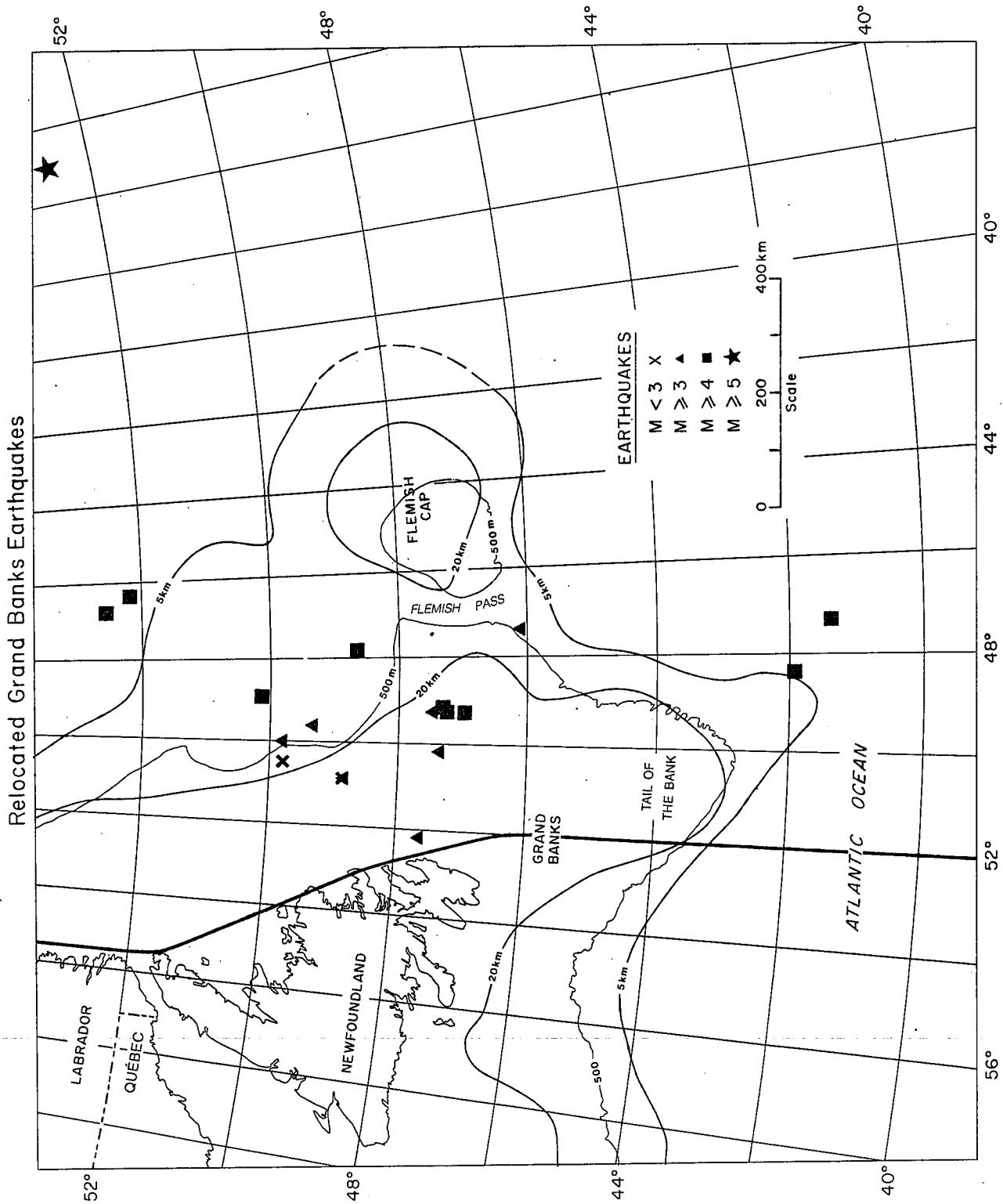


Figure 10. Relocated offshore earthquakes, crustal thickness (taken from Shih et al., 1988), and 500-m isobath. Compare with Figure 3 to determine the identity of the events. As in previous maps, earthquakes west of the heavy line are not plotted.



APPENDIX

Event files in this Appendix are arranged chronologically. We have tried to provide a full documentation of our data and results, so for each earthquake there is a "PIK" listing of the earthquake phase data used to compute the epicentre (a description of the format is given on the final pages), together with comments on the data and epicentre.

22

```
+70.000- 75.0000LMS=5.1 0628550 26071922 00.0000.000 0.0 2 3 00.00N218.00 0 1ML=0.0 00 0L3.60
$50.0 - 50.0 ML=5.3 063108 CEEF
BAFFIN ISLAND REGION
$ ORIGINALLY PLACED BY ISS AT 50N 50W
$ EPICENTRE PEGGED AT LOCATION FOUND ROUGHLY BY WAHLSTROM
$ SEE GSC OPEN FILE 3043 BY ADAMS AND WAHLSTROM
$ DOUBTFUL THAT ANY BETTER CAN BE DONE WITHOUT THE ORIGINAL RECORDS
$ SAS OPERATING BUT RECORD NOT KEPT
$ HAL READINGS BY WAHLSTROM FORCES EPICENTRE TO NORTH
$ CHI, WAS, ESK READINGS FROM ISS, REPRESENT RAYLEIGH WAVE MAXIMUM
$ SEE TEXT OF GSC OPEN FILE 3043 FOR DISCUSSION
$ OTT P TAKEN WITH ASSUMED 1 MIN READING OR REPORTING BLUNDER FITS HAL S IF PHASE IS FIRST S
$ LOCATION FROM HAL P AND SN, OTT SN AND (P-1MIN) IS:
$68.837- 73.7630LMS=5.1 0629094 26071922 00.3742.100 0.0 3 4 02.52 218.00 0 1ML=0.0 00 0 3.62
$ GSC DOES NOT CURRENTLY HAVE A TELESEISMIC LOCATION PROGRAM TO ALLOW
$ FITTING THE DISTANT PHASES
$
OTT 2207260631P X3514 3837 X4006
OTT S 2744KM 00 6012$181 35 00-1148$00-9123$ 0000000 00ML00MN
HAL 2207260631P 3433 3905 X3905
HAL S 2899KM 00 622 161 34 00-1646$00-9999$ 0000000 00ML00MN
CHI 2207260631P X4215
CHI S 3215KM 200 33 00-9307$ 0000000 00ML00MN
WAS 2207260631P X4400
WAS S 3468KM 183 33 00-5841$ 0000000 00ML00MN
ESK 2207260631P X4400
ESK E 3748KM 078 32 00-9999$ 0000000 00ML00MN
Z
```

62

```
+61.027- 58.2250 ML=4.8 0131042 03081962 00.0000.000 0.0 2 3 10.00 218.00 0 1MN=0.0 00 0 3.65
$+52 -54.2 ML=4.8 CEEF AFTER SMITH
$ LABRADOR RIDGE (PROBABLY)
$ NEEDS GREENLAND STATIONS AND RES, MBC, ALE, PNT.
$
$ SMITH'S COMMENTS
$ SCH CLEAR PN AND SN; NO LG - OCEANIC EVENT
$ HAL (SN) TRACE "VERY SMALL AND MAY NOT BE ASSOCIATED".
$ WESTERN EPICENTRE NOT POSSIBLE BECAUSE OF LACK OF RECORDS ON SOUTHERN STATIONS.
$
$ WAHLSTROM COMMENTS:
$ HAL: RE-READ AT LAMONT, LARGE T.C. USED TO REPORT PHASE IN BULLETIN
$ CONFIRMED WITH TELESEISM, CHECK AGAIN FOR S.
$ SCH: ARR. TIMES AND AMP/PER FROM SMITH CONFIRMED (ROUGHLY) BY MICROFILM,
$ RECORD NOT FOUND? (CHECK)
$ SFA: MICROFILM NOT FOUND, RECORD NOT FOUND, INTERMITTENT OPERATION
$ SHF: NO SIGNAL FROM MICROFILM, RECORD NOT FOUND? (CHECK)
$ OTT: STATION DISCONNECTED.
$ MNT: POSSIBLE S PHASE AT 3904 ON EW, NOT ON NS. 44 SEC TOO LATE FOR SN
$
$ ADAMS COMMENTS
$ NO TRACE OF SECOND (EARLIER OR LATER) PHASE ON HAL
$ HAL PHASE RE-INTERPRETED AS PN, PLACES EPICENTRE ON LABRADOR RIDGE.
$ CURRENT EXPERIENCE SUGGESTS SMITH EPICENTRE JUST N OF NEWFOUNDLAND
$ SHOULD HAVE GENERATED LG ON SCH
```

\$ RIDGE LOCATION IS CONSISTANT WITH OCEANIC PATH FOR SCH.
 \$ RIDGE IS LOCUS OF MANY EARTHQUAKES; SMITH'S LOCATION IS HIGHLY UNUSUAL.
 \$ ABSENCE OF SN ON HAL IS STILL PUZZLING.
 \$ COMPARE WITH ML=5.0 EVENT ON 621202.
 \$ COMPARE WITH ML=5.1 EVENT ON 621026. THIS EVENT ALSO LACKS HAL SN READING
 \$ THOUGH AS IT WAS LARGER HAD PN ON SHF AND RES TO CONFIRM NORTHERN EPICENTRE.
 \$

| | | | | | | | |
|-----|--------------|------------|--|--------|--|------------------|-----|
| SCH | 6208030132P | 32545 | | 3416 | | 30 175 150 | 8 1 |
| SCH | SW 0857KM 10 | 000 220 49 | | 10 000 | | 0001795 48ML42MN | |
| HAL | 6208030132P | 3456 | | | | | 1 |
| HAL | S 1860KM 10 | 000 193 50 | | | | 0000000 00ML00MN | |

69

+47.632- 52.156FLMN=3.4 2153207 05081969 00.0860.110 0.2 5 9 40.94 218.00 0 1ML=0.0 00 0 3.62
 \$47.74 - 52.31 MN=3.3 215324 CEEF

\$ OFF ST JOHN'S
 \$ ST JOHN'S NEWSPAPERS DO NOT REPORT IT AS FELT, DESPITE CLOSENESS
 \$ COMMENT FROM COMPUTER CARDS: FM IS CNW - EVENT FROM ALMOST DUE EAST
 \$ HAL: LG USED FOR MAGNITUDE, MAY BE ATTENUATED
 \$

| | | | | | | | |
|-----|--------------|-------------|----------|--------|---------|------------------|---|
| STJ | 6908052153P | | A53288 C | | A53333 | | 1 |
| STJ | W 0044KM | 261-68 | 03 036 | | 08 -059 | 0000000 00ML00MN | |
| HAL | 6908052153P | B5523 | | B5652 | X5746 | 040 66 6 | 1 |
| HAL | W 0944KM 03 | 142 254 49 | | 01 089 | 00 446 | 0000143 40ML32MN | |
| SIC | 6908052153P | A55421 | | | B58275 | 050 106 10 | 1 |
| SIC | W 1105KM 18 | 088 290 49 | | | 03 154 | 0000119 42ML32MN | |
| SCH | 6908052153P | A56032 | | | B5922 | 050 72 11 | 1 |
| SCH | NW 1294KM 32 | -117 314 49 | | | 20 365 | 0000192 46ML36MN | |
| SFA | 6908052153P | B56190 | | | | | 1 |
| SFA | W 1409KM 01 | 071 275 49 | | | | 0000000 00ML00MN | |
| GWC | 6908052153P | | | | XB6224 | 090 35 4 | 1 |
| GWC | NW 1959KM | 305 47 | | | 00 195 | 0000080 47ML35MN | |

70

+52.189- 46.33301ML=4.3 1744440 31101970 00.0310.110 0.3 2 4 20.29 218.00 0 1MN=0.0 00 0 3.62
 \$ 52.17 46.37 ML=4.3 174445 CEEF

\$ GRAND BANKS
 \$ ORIGINAL COMMENTS: NO RECORDS FOR SIC; NOT AT HAL, FBC, SFA, GWC
 \$ WAHLSTROM COMMENTS: NOT ON HAL, OTT, GWC, FBC, SFA, SUD
 \$ SIC CLOCK WRONG
 \$ NO SIGNAL ON HAL EXCLUDES THE OTHER POSSIBLE SOLUTION: 43.3N 59.3W
 \$ ML ASSUMED - EARTHQUAKES CLOSE BY DO NOT HAVE LG
 \$

| | | | | | | | |
|-----|--------------|-------------|--|---------|--|------------------|-----|
| STJ | 7010311744P | 46135 | | 47205 | | 30 34 9 | 8 1 |
| STJ | SW 0689KM 15 | -036 224 49 | | 05 021 | | 0000554 41ML36MN | |
| SCH | 7010311744P | 4739 | | 4948 | | 40 94 9 | 8 1 |
| SCH | W 1384KM 15 | 036 290 49 | | 05 -021 | | 0000150 45ML35MN | |

71

+47.308- 49.160FLMN=4.8 0617095 15081971 00.1310.215 0.2 10 17 62.50 218.00 0 1ML=0.0 00 013.62
 \$ 47.46- 49.53 MB=4.8 061715 CEEF (FROM ISC)
 \$ 47.46- 49.53 MB=4.8 0617152 ISC

\$ GRAND BANKS
 \$ READING BY WAHLSTROM IN 1985 ADDED BY ADAMS OCT 1988
 \$ POLARITIES CHECKED BY ADAMS OCT 1988
 \$ WAHLSTROM COMMENTS: STJ ONLY GIVES DIST=255KM, BACKAZ=084DEG, H=18KM ASSUMED
 \$ I.E. 47.8N 49.4W OT=061713
 \$ STJ: FM IS D-E. EVENT FROM SLIGHTLY NORTH OF EAST
 \$ WELL-RECORDED ON LPS
 \$ HAL: FM IS C. LG SLIGHTLY LARGER THAN SN
 \$ SCH READ BY WAHLSTROM AS C; BY ADAMS AS -
 \$ SIC READ BY WAHLSTROM AS C; BY ADAMS AS +
 \$ OTT READ BY WAHLSTROM AS ; BY ADAMS AS +
 \$

\$ CHECK ISC
 \$ CHECK REST OF CSN
 \$ WAHLSTROM HAS TRIED A FM
 \$

| | | | | | | | |
|-----|-------------|------------|---------|----------------|--------|------------------|---|
| STJ | 7108150617P | A17498 D | X17508 | B18172 | X18192 | | 1 |
| STJ | W 0271KM 05 | 147 278 49 | 00 -259 | 00 042 00 -544 | | 0000000 00ML00MN | |

| | | | | | | | | | |
|-----|-------------|--------|----|--------|------|----|---------|----------|----------|
| HAL | 7108150617P | A19371 | C | A21278 | 100 | 18 | 155 | 3 | 1 |
| HAL | W | 1156KM | 02 | 087 | 260 | 49 | | | |
| SIC | 7108150617P | A19560 | C | 17 | 277 | | 0005411 | 60ML49MN | |
| SIC | W | 1329KM | 04 | -137 | 290 | 49 | | | |
| UNB | 7108150617P | A19533 | - | | | | 0000000 | 00ML00MN | |
| UNB | W | 1346KM | 81 | -604 | 270 | 49 | XA23336 | 100 | 35 |
| SCH | 7108150617P | A20167 | C | | | | 00 | 1232\$ | 0002980 |
| SCH | NW | 1486KM | 00 | 015 | 311 | 49 | A22344 | 80 | 62 |
| SFA | 7108150617P | A20330 | D | 02 | -098 | | | | 0005130 |
| SFA | W | 1637KM | 08 | -193 | 277 | 49 | A23069 | X25001 | 90 |
| OTT | 7108150617P | A21213 | | 01 | -054 | 00 | 1826\$ | 0001164 | 57ML45MN |
| OTT | W | 2047KM | 01 | -074 | 274 | 47 | B24410 | 110 | 17 |
| GWC | 7108150617P | A21376 | + | 06 | 650 | | | | 0000941 |
| GWC | NW | 2165KM | 21 | 305 | 305 | 45 | B24578 | 100 | 26 |
| FBC | 7108150617P | B21389 | | 01 | -194 | | | | 0002779 |
| FBC | NW | 2181KM | 01 | 269 | 334 | 43 | B25114 | 90 | 39 |
| SUD | 7108150617P | A22011 | + | 09 | 827 | | | | 0000322 |
| SUD | W | 2411KM | 12 | 229 | 280 | 39 | | | 110 |
| | | | | | | | | | 20 |
| | | | | | | | | | 49 |
| | | | | | | | | | 0001399 |
| | | | | | | | | | 61ML49MN |

76

+50.133- 48.910F1ML=4.1 1923304 28081976 00.0260.108 0.3 5 8 30.54 218.00 0 1MN=0.0 00 0 3.62
 \$50.10 - 48.85 ML=4.0 192330 CEEF
 \$ GRAND BANKS
 \$ MUNF IS MEMORIAL SEISMOGRAPH IN STJ VAULT
 \$ WAHLSTROM COMMENTS:
 \$ ISC: AS OTT SOLUTION, I.E., 50.10 N 48.85 W OT=192330 ML=4.0 "POOR SOLUTION"
 \$ MIQ, SUD, FRB, OTT, QCO, LGQ: NOT RECORDED, CBK, HAL: RECORD NOT FOUND
 \$ PBQ: VISIBLE 1932 BUT NO CLEAR ONSET
 \$ UNB, SIC: PROB DISTURBANCES AT APPR 2940
 \$

| | | | | | | | | | | | |
|------|-------------|--------|----|------|-----|----|--------|------|---------|----------|----------|
| MUNF | 7608281925P | | | | | | 40 | 73 | 156 | 8 | 1 |
| MUNF | SW | 0400KM | | 226 | 49 | | | | 0003357 | 42ML39MN | |
| STJ | 7608281925P | A24258 | D | | | | A25048 | 40 | 32 | 20 | 8 |
| STJ | SW | 0400KM | 34 | 086 | 226 | 49 | 04 | -028 | | 0000982 | 37ML34MN |
| SCH | 7608281925P | A26168 | | | | | A28209 | 30 | 100 | 8 | 8 |
| SCH | NW | 1318KM | 01 | -017 | 300 | 49 | 06 | 037 | | 0000168 | 43ML35MN |
| MNQ | 7608281925P | A26281 | | | | | B28427 | 50 | 5 | | 3 |
| MNQ | W | 1412KM | 04 | -031 | 279 | 49 | 14 | 219 | | 0000000 | 00ML00MN |
| POC | 7608281925P | A26483 | | | | | | | | | 1 |
| POC | W | 1580KM | 12 | -052 | 267 | 49 | | | | 0000000 | 00ML00MN |
| CHQ | 7608281925P | B27006 | | | | | | | | | 1 |
| CHQ | W | 1688KM | 05 | -137 | 266 | 49 | | | | 0000000 | 00ML00MN |

77

+49.337- 49.58601ML=3.4 0805261 25031977 00.1230.155 0.0 2 4 10.12 218.00 0 1MN=0.0 00 0 3.62
 \$ NEW EVENT FOUND BY SEARCHING MEMORIAL UNIVERSITY SEISMOGRAMS
 \$ GRAND BANKS
 \$ ALTERNATIVE LOCATION (LESS LIKELY BECAUSE HAL SHOWS NO TRACE) FOLLOWS:
 \$44.88 - 52.59 01ML=3.4 080527. 25031977 1 0.06 0.54 0.0 2 5 1 0.5 218MUN 300
 \$ CBK PEN AND INK RECORD
 \$ MUNF NO LG
 \$

| | | | | | | | | | | | | | |
|------|-------------|--------|----|------|-----|----|---------|--------|-----|-------|------|---------|----------|
| MUNF | 7703250806P | A0609 | | | | | XB06175 | A06405 | 030 | 108 | 85 | 8 | 1 |
| MUNF | SW | 0305KM | 00 | -001 | 231 | 49 | 00 | 213 | 00 | 001 | | 0001648 | 34ML34MN |
| CBK | 7703250806P | C0649 | | | | | | | | C0745 | | | 1 |
| CBK | W | 0614KM | 30 | 227 | 269 | 49 | | | | 10 | -130 | | 0000000 |
| HAL | 7703250806P | | | | | | | | | | | | 00ML00MN |
| HAL | W | 1185KM | | 249 | 49 | | | | | | | | 0000000 |
| | | | | | | | | | | | | | 00ML00MN |

78

+52.573- 46.752F1ML=4.1 1435140 03011978 00.0280.073 0.1 3 6 30.36 218.00 0 1MN=0.0 00 0 3.62
 \$ NEW EVENT FOUND BY SEARCHING MEMORIAL UNIVERSITY SEISMOGRAMS
 \$ GRAND BANKS
 \$ ADAMS RE-READ RECORDS OCT 1988
 \$ MUNF, STJ NO LG
 \$

| | | | | | | | | | | | | | |
|------|-------------|---------|------|------|-----|----|--------|-----|-----|-----|----|---------|----------|
| MUNF | 7801031436P | XA36458 | - | | | | X37542 | | 030 | 108 | 30 | 8 | 1 |
| MUNF | SW | 0702KM | 00 | 038 | 220 | 49 | 00 | 120 | | | | 0000582 | 41ML36MN |
| STJ | 7801031436P | A36455 | - | | | | A37530 | | 040 | 29 | 6 | 8 | 1 |
| STJ | SW | 0702KM | 01 | 008 | 220 | 49 | 00 | 000 | | | | 0000325 | 40ML33MN |
| CBK | 7801031436P | A3710 | -2.0 | | | | A38350 | | | | | | 1 |
| CBK | SW | 0889KM | 08 | -029 | 247 | 49 | 01 | 010 | | | | 0000000 | 00ML00MN |

| | | | | | |
|-----|-------------|---------------|---------|------------------|-----|
| SCH | 7801031436P | B38066 | A40094 | 050 120 7 | 8 1 |
| SCH | W 1343KM | 50 295 289 49 | 01 -008 | 0000073 42ML32MN | |
| FRB | 7801031436P | | | | 1 |
| FRB | NW 1773KM | 323 49 | | 0000000 00ML00MN | |
| | Z | | | | |

81

+48.643- 47.820F1ML=4.2 1306126 26091981 00.0670.170 0.2 5 10 61.16 218.00 0 1MN=0.0 00 0 3.62
 \$ NEW EVENT FOUND BY SEARCHING MEMORIAL UNIVERSITY SEISMOGRAMS
 \$ GRAND BANKS
 \$ READ BY WAHLSTROM AND ADAMS
 \$MUNF: TC UNKNOWN, APPEARS TO BE 0.0 AS TIMES AGREE
 \$MUNF: LG=0.5SN, NO AFTERSHOCKS
 \$CBK: T.C. UNKNOWN, T.C. -2 SEC APPLIED
 \$ LG WEAK OR ABSENT
 \$HAL: V QUIET TRACE
 \$SIC: QUIET, NO PHASES READABLE
 \$UNB: BOTH READING SIMILAR TO NOISE
 \$LPQ,LMQ: HIGH FREQUENCY ON LOW FREQUENCY NOISE
 \$FRB: NOT RECORDED
 \$

| | | | | | | |
|------|-------------|----------------|--------|---------|------------------|-----|
| STJ | 8109261307P | A07068 | A07155 | A07431 | 50 36 65 | 8 1 |
| STJ | W 0385KM | 24 153 254 49 | 05 073 | 09 -095 | 0002269 41ML37MN | |
| MUNF | 8109261307P | | | | 30 108 250 | 8 1 |
| MUNF | W 0385KM | 254 49 | | | 0004848 42ML41MN | |
| CBK | 8109261307P | X07523 | | X09031 | 50 89 | 3 1 |
| CBK | W 0746KM | 00 291 276 49 | | 00 208 | 0000000 00ML00MN | |
| HAL | 8109261307P | X09046 | | B10566 | 50 83 6 | 8 1 |
| HAL | W 1285KM | 00 957 256 49 | | 01 106 | 0000091 43ML32MN | |
| UNB | 8109261307P | XB09168 | | | | 1 |
| UNB | W 1452KM | 00 139 265 49 | | | 0000000 00ML00MN | |
| SCH | 8109261307P | A09180 | | A11369 | 40 95 6 | 8 1 |
| SCH | NW 1474KM | 00 -016 305 49 | | 10 100 | 0000099 44ML34MN | |
| MNQ | 8109261307P | A09224 | | B11437 | 40 220 8 | 8 1 |
| MNQ | W 1526KM | 43 -204 286 49 | | 06 -315 | 0000057 42ML31MN | |
| LPQ | 8109261307P | XA09410 | | | | 1 |
| LPQ | W 1658KM | 00 050 273 49 | | | 0000000 00ML00MN | |
| LMQ | 8109261307P | A09427 | | B12200 | 30 190 3 | 8 1 |
| LMQ | W 1676KM | 00 -006 274 49 | | 01 119 | 0000033 39ML30MN | |
| | Z | | | | | |

82

+41.367- 47.282F1ML=4.0 0407239 22071982 00.1040.124 0.1 6 8 41.05 218.00 0 1MN=0.0 00 0 3.62
 \$ NEW EVENT FOUND BY SEARCHING MEMORIAL UNIVERSITY SEISMOGRAMS
 \$ SOUTHEAST OF TAIL OF THE BANK, GRAND BANKS
 \$ FIRST EARTHQUAKE KNOWN IN THIS AREA
 \$ GOOD ON MUNF, STJ AND CBK, KLN
 \$ HAL: VERY WEAK
 \$ CBK TC SEVERAL HOURS AND ABOUT 26 SECONDS
 \$ SCH, UNB: PRESENT, BUT UNCLEAR
 \$ SIC, GSQ, LMN, GGN, LMQ: RECORDS NOT CHECKED
 \$

| | | | | | | |
|------|-------------|-----------------|---------|--|------------------|-----|
| STJ | 8207220409P | X09114 | A10267 | | 50 36 4 | 8 1 |
| STJ | NW 0815KM | 00 239 330 49 | 00 -009 | | 0000140 39ML31MN | |
| MUNF | 8207220409P | A09087 | X10277 | | 20 173 19 | 8 1 |
| MUNF | NW 0815KM | 01 -031 330 49 | 00 090 | | 0000345 39ML35MN | |
| CBK | 8207220409P | X1021 | X12123 | | 50 11 | 3 1 |
| CBK | NW 1186KM | 00 022 319 49 | 00 -006 | | 0000000 00ML00MN | |
| HAL | 8207220409P | XB10284 | B12265 | | | 1 |
| HAL | W 1377KM | 00 1084\$291 49 | 00 011 | | 0000000 00ML00MN | |
| KLN | 8207220409P | A10491 | A13210 | | 20 628 18 | 8 1 |
| KLN | NW 1641KM | 05 -070 298 49 | 28 -164 | | 0000090 41ML34MN | |
| EBN | 8207220409P | A11082 | A13573 | | | 1 |
| EBN | NW 1795KM | 00 -018 299 52 | 32 177 | | 0000000 00ML00MN | |
| MNQ | 8207220409P | A11264 | XC14327 | | 50 161 2 | 8 1 |
| MNQ | NW 1944KM | 14 115 309 50 | 00 557 | | 0000016 39ML28MN | |
| | Z | | | | | |

84a

+48.852- 50.83701MN=3.6 1925475 26051984 00.0190.037 0.3 8 18 60.48 218.00 0 1ML=4.1 30 0 3.62
 \$48.89 - 51.06 MN=3.5 192540 CEEF
 \$ GRAND BANKS
 \$ KLN,SIC NOISY.
 \$ LG>SN FOR THIS OFFSHORE EVENT
 \$ AFTERSHOCK ONLY ON MUNF AT 2204:50, SAME S-P, MAGNITUDE = 2.6

\$
 STJ 8405261926P A26181 A26395 C26405 040 29 167 1
 STJ SW 0200KM 18 038 225 49 07 -024 04 -266 0009046 37ML39MN
 CBK 8405261926P A26566 A27490 030 184 125 1
 CBK W 0523KM 35 -053 274 49 34 051 0001423 41ML38MN
 GBN 8405261926P A27427 A29075 040 211 35 1
 GBN W 0896KM 05 021 249 49 03 -014 0000261 42ML34MN
 HAL 8405261926P C2805 X3002 C3049 1
 HAL W 1080KM 00 -003 249 49 00 1504\$04 295 0000000 00ML00MN
 SCH 8405261926P B2829 C3027 C3138 05 102 15 1
 SCH NW 1281KM 02 -054 307 49 04 -273 06 -347 0000185 46ML35MN
 GGN 8405261926P C2830 B3029 04 132 14 8 1
 GGN W 1283KM 00 026 257 49 09 -107 0000167 44ML35MN
 MNQ 8405261926P C2831 C3030 C31415 04 163 20 1
 MNQ W 1305KM 01 -154 285 49 12 -495 23 -674 0000193 45ML36MN
 JAQ 8405261926P C3403 05 163 06 1
 JAQ NW 1811KM 297 52 11 -479 0000046 43ML32MN
 Z

84b

+48.852- 50.8370LMN=2.6 2204198 26051984 00.0000.000 0.0 1 1 00.00H218.00 0 1ML=0.0 00 0 3.62
 \$ NEW EVENT NOT IN CEEF
 \$ GRAND BANKS
 \$ AFTERSHOCK OF EARTHQUAKE AT 192547
 \$ ONLY ON MUNF (STJ) AT 2204:50, SAME S-P, MAGNITUDE = 2.6 BY PROPORTION TO MAINSHOCK
 \$ PEGGED AT MAINSHOCK EPICENTRE
 \$
 STJ 8405262204P A0450
 STJ SW 0200KM 00 000 225 49 0000000 00ML00MN
 Z

85

+49.787- 50.480F1ML=2.8 1806035 03081985 00.0150.053 0.3 4 6 20.14 218.00 0 1MN=2.8 10 0 3.62
 \$ NEW EVENT FOUND BY SEARCHING MEMORIAL UNIVERSITY SEISMOGRAMS
 \$ WELL RECORDED ONLY ON STJ AND MUNF
 \$ HAL, SCH, MNQ, SIC, KLN, LPQ, LMQ, NOTHING
 \$ SLQ AT PAPER CHANGE
 \$ POOR RECORDING ON GBN CONFIRMS THIS EVENT IS NE OF STJ, AND NOT A LSP EQ
 \$ MUNF PN NEAR PAPER EDGE
 \$ HINT OF LG ON MUNF
 \$ GBN GOOD PHASE MUST BE PN TO FIT STJ ORIGIN TIME
 \$ CBK READINGS VERY POOR, MAY BE FOLLOWED BY LG AT 0841
 \$ SN ABSENT ON GGN, TINY ON GBN (TYPICAL FOR THESE EQ)
 \$ NOT A WELL-LOCATED EARTHQUAKE, BUT CLEARLY ON NORTHERN GRAND BANKS
 \$
 MUNF 8508031806P XB06455 XA07155 020 173 22
 MUNF SW 0297KM 00 001 215 49 00 -080 0000400 26ML28MN
 STJ 8508031806P A06455 A07163 040 29 11 8
 STJ SW 0297KM 00 001 215 49 00 -001 0000596 30ML30MN
 CBK 8508031806P C0713 C0812
 CBK W 0553KM 53 -369 263 49 06 128 0000000 00ML00MN
 GBN 8508031806P A08065 XC09435
 GBN SW 0961KM 00 001 244 49 00 588 0000000 00ML00MN
 GGN 8508031806P C0852
 GGN W 1335KM 00 -002 253 49 0000000 00ML00MN
 Z

88a

+49.834- 49.982F1ML=3.5 0715448 09011988 00.0180.058 0.2 8 11 60.34 218.00 0 1MN=0.0 00 0 3.62
 \$49.880- 49.883 ML=3.5 0715438 CEEF
 \$ ON NORTHERN SIDE OF GRAND BANKS; AN AREA OF FEW EARTHQUAKES
 \$ SOLUTION BY ADAMS, WITH A GREAT DEAL OF EFFORT
 \$ HAL: NO SN VISIBLE
 \$ KLN MAX AMPLITUDE TINY
 \$
 MUNF 8801090715P A16297 XA17045 30 91 70 8
 MUNF SW 0323KM 11 -022 220 49 00 142 0001611 35ML35MN
 STJ 8801090715P A17030 40 29 15 8
 STJ SW 0323KM 220 49 02 -008 0000812 33ML32MN
 CBK 8801090715P A17023 B17595 40 135 21 8
 CBK W 0589KM 01 -008 263 49 01 -020 0000244 35ML31MN
 GBN 8801090715P A17525 B1926 20 563 13 8
 GBN SW 0996KM 59 052 245 49 01 -024 0000073 35ML29MN
 HAL 8801090715P C1817
 HAL SW 1180KM 06 261 246 49 0000000 00ML00MN

EEO 8905142257P B593575 -0.06
EEO W 3200KM 00 -035 276 33
Z

000 0 0 0 0
0000000 00ML00MN

89b

+46.960- 49.247F1ML=4.2 0515398 03121989 00.0500.088 0.4 6 9 40.71 218.00 0 1MN=0.0 00 0L3.62
OFFSHORE NEWFOUNDLAND AU LARGE DE TERRE-NEUVE

MAG (GSC) 3.9 MB (1 OBS)
\$ MB 3.9 FROM YELLOWKNIFE ARRAY
\$ MNQ DOWN, TOO WEAK FOR OTHERS
\$ LOCATION FROM YELLOWKNIFE ARRAY: 49.7 51.4 O.T. 05:15:58
\$ SCH LG>=SN AMP; SN PRECEDED BY NOISE
\$ STJ UNABLE TO DECIDE IF LG>SN
\$ FRB NOT VISIBLE
\$ CBK NOT OPERATING
\$ ADAMS ADDED STJ AND REVISED SCH
\$ ADAMS DID NOT USE KLN, GGN, EBN BECAUSE TOO MANY STATIONS ON THIS AZIMUTH

\$
STJ 8912030515P B16197 XB16215 C1647
STJ W 0273KM 19 092 286 49 00 -239 00 -035 0000000 00ML00MN
GBN 8912030515P B17435 - B19145 040 216 20 8
GBN W 0962KM 09 062 264 49 04 043 0000145 40ML32MN
HAL 8912030515P B1804 0000000 00ML00MN
HAL W 1144KM 22 -100 262 49 000 0 0 0 0
LMN 8912030515P B181142 -0.29 XC2005 000 0 0 0 0
LMN W 1201KM 20 -096 270 49 00 -034 0000000 00ML00MN
KLN 8912030515P XB182287 -0.29 XC2031 050 204 4 0 8
KLN W 1303KM 00 -192 276 49 00 402 0000025 37ML27MN
GGN 8912030515P XB183188 -0.29 XC2043 050 102 7 0 8
GGN W 1374KM 00 -147 268 49 00 109 0000086 43ML32MN
EBN 8912030515P XB1838 0.00 000 0 0 0 0
EBN W 1437KM 00 -285 279 49 0000000 00ML00MN
SCH 8912030515P B18500 - B21095 050 102 11 8
SCH NW 1507KM 09 065 312 49 06 -054 0000136 46ML35MN
YKA 8912030515P B2304 0.00 000 0 0 0 0
YKA NW 4329KM 00 -010 318 31 0000000 00ML00MN
Z

90

+41.918- 48.34101ML=4.7 2140041 24041990 00.0270.062 0.1 14 23 130.67 218.00 0 1MN=0.0 00 0L3.62
042.000- 48.48001MB=4.3 2140030 24041990 00.1800.072 0.1 00 00 81.31 618.00 0 1MN=0.0 00 0L3.62
SOUTHERNMOST PART OF THE GRAND BANKS LA PARTIE LA PLUS AU SUD DU GRAND BANC
DE TERRE NEUVE

MAG (GSC) 3.9 MB (1 OBS)
MAG (ISC) 4.3 MB (8 OBS)
\$ DETECTED ON THE YELLOWKNIFE ARRAY
\$ MAGNITUDE MB=3.9 AT 1 HZ ON YK
\$ STJ: NO LG; POLARITY POSSIBLE +
\$ SCH: STRONG P, WEAK SN; PN READING IS VERY SHARP
\$ GBN: EMERGENT BEGINNING, POLARITY NOT READABLE
\$ CBK: PRECURSOR TO SN MAKES ONSET UNCERTAIN

\$
STJ 9004242140P A41379 A42465 020 173 145 8
STJ NW 0718KM 27 049 333 49 00 005 0002633 46ML43MN
CBK 9004242140P B4221 XB4405 040 132 46 8
CBK NW 1082KM 05 -082 319 49 00 106 0000547 48ML39MN
GBN 9004242140P B42295 E B4412 020 563 200 8
GBN NW 1130KM 25 190 295 49 28 -203 0001116 48ML42MN
HAL 9004242140P B42455 XC44405 030 207 60 8
HAL W 1273KM 01 047 289 49 00 -394 0000607 49ML40MN
LMN 9004242140P A425987D -0.29 B451097 013 100 13 0 8
LMN NW 1391KM 01 008 294 49 07 100 0000628 46ML41MN
GGN 9004242140P C432033 -0.29 C453877 017 100 13 0 8
GGN W 1533KM 05 328 290 49 01 -131 0000480 47ML41MN
UNB 9004242140P C43175 X4532 050 74 12 8
UNB NW 1533KM 00 072 293 49 00 -783 0000204 48ML37MN
KLN 9004242140P B431925 -0.29 C454049 0131000 88 0 8
KLN NW 1534KM 28 202 297 49 00 010 0000425 46ML40MN
GSQ 9004242140P B433283 -0.22 C460650 040 100 14 0 8
GSQ NW 1657KM 03 069 304 49 00 004 0000220 48ML38MN
EBN 9004242140P B433497 -0.22 C461354 0201000 77 0 8
EBN NW 1689KM 07 -101 298 49 00 038 0000242 46ML38MN
LPQ 9004242140P XB435397 -0.22 000 0 0 0 0
LPQ NW 1816KM 00 289 297 52 0000000 00ML00MN
DAQ 9004242140P B440216 -0.06 C470492 0271000 83 0 8
DAQ NW 1921KM 04 -077 298 50 02 241 0000193 48ML38MN

| | | | | | | | | |
|------|-------------|-----------|----------|---------|---------|----------|----|---|
| SCH | 9004242140P | A44076 | C | C4713 | 060 | 94 | 12 | 8 |
| SCH | NW 1971KM | 74 | -082 323 | 00 -003 | 0000134 | 49ML37MN | | |
| DPQ | 9004242140P | C441495 | -0.06 | X472281 | 0331000 | 46 | 0 | 8 |
| DPQ | NW 2013KM | 02 | 194 294 | 00 089 | 0000088 | 45ML35MN | | |
| YKB9 | 9004242140P | XB480459C | | | 0000000 | 00ML00MN | | |
| YKB9 | NW 4803KM | 00 | 117 322 | | 000 | 0 | 0 | 0 |
| YKA | 9004242140P | X4804 | 0.00 | | 0000000 | 00ML00MN | | |
| YKA | NW 4804KM | 00 | 053 322 | | | | | |
| YKR3 | 9004242140P | XB480555C | | | 0000000 | 00ML00MN | | |
| YKR3 | NW 4816KM | 00 | 118 322 | | | | | |
| YKR2 | 9004242140P | B480575C | | | 0000000 | 00ML00MN | | |
| YKR2 | NW 4819KM | 10 | 119 322 | | | | | |

Z

91

+47.364- 50.167F1MN=3.2 1123013 23071991 00.0820.146 0.2 6 10 30.79 218.00 0 1ML=3.8 20 0L3.62
 GRAND BANKS

\$ EAST OF ST JOHN'S
 \$ LG ON STJ, CBK, SMQ, SCH, NOT ON GBN OR LMN
 \$ NOT ON LMQ, CIQ, DAQ
 \$ MNQ DEAD
 \$ NO ECTN TRIGGER (?)
 \$ STJ: ONLY FIRST MOTION VISIBLE; NO AFTERSHOCKS EVIDENT
 \$ CBK: P PRECEDED BY ?NOISE AT 24122
 \$ CBK: NO CLEAR S PHASES
 \$ HAL: NOISY RECORD, READING IS PROBABLY JUST NOISE
 \$ SMQ: NOISY RECORD
 \$ SCH: LG SEEMS OK
 \$ KUQ: ?PHASE AT 26455 IS 10 SEC TOO LATE FOR P
 \$
 \$ SOLUTION BY ADAMS 910814
 \$ POOR STATION DISTRIBUTION; THIS LOCATION FITS PN'S AT STJ, CBK, GBN, AND SCH
 \$ ACCEPTABLY, ALSO SN'S AT GBN AND SCH (BUT NOT CBK S PHASES)
 \$

| | | | | | | | | |
|-----|-------------|---------|----------|-----------|-----------|----------|----------|----|
| STJ | 9107231123P | A23308 | D | | | | | |
| STJ | W 0195KM | 00 | -005 278 | | 0000000 | 00ML00MN | | |
| CBK | 9107231123P | B2420 | | C2515 | XC2535 | 030 | 193 | 27 |
| CBK | W 0606KM | 07 | -100 289 | 10 -486 | 00-1382\$ | 0000293 | 37ML32MN | |
| GBN | 9107231123P | B24575 | | B26205 | | 020 | 563 | 25 |
| GBN | W 0899KM | 04 | 077 260 | 20 -171 | | 0000140 | 36ML31MN | 8 |
| HAL | 9107231123P | | | X2715 | | | | |
| HAL | W 1082KM | | 259 49 | 00 1383\$ | | 0000000 | 00ML00MN | |
| LMN | 9107231123P | B2525 | | B27145 | | 040 | 300 | 13 |
| LMN | W 1133KM | 01 | -028 267 | 43 248 | | 0000068 | 39ML30MN | 8 |
| SMQ | 9107231123P | XB25325 | | C2734 | X2832 | 040 | 300 | 13 |
| SMQ | W 1254KM | 00 | -752 291 | 06 -373 | 00-1579\$ | 0000068 | 40ML31MN | |
| SCH | 9107231123P | B2602 | | B2814 | XB29155 | 060 | 94 | 10 |
| SCH | NW 1425KM | 09 | 112 312 | 00 -013 | 00-1950\$ | 0000111 | 45ML34MN | |

Z

92a

+47.237- 49.226F1MN=4.0 0607280 13011992 00.0560.097 0.2 16 24 140.98 218.00 0 1ML=4.5 20 0L3.62
 SOUTHERN GRAND BANKS

OFFSHORE NEWFOUNDLAND AU LARGE DE TERRE-NEUVE
 \$ EXACT EPICENTRE RATHER UNCERTAIN BECAUSE OF EXTREMELY POOR STATION DISTRIBUTION
 \$
 \$ ISC 47.2N 49.2W OT=060727.2; added 3 stations to GSC's 15;
 \$ "EXACT EPICENTRE RATHER UNCERTAIN BECAUSE OF EXTREMELY POOR STATION DISTRIBUTION"
 \$ ADAMS SOLUTION
 \$ VLG OF 3.62 USED FOR THIS SOLUTION
 \$ LG GREATER THAN SN ON MOST ECTN EXCEPT LMN AND DAQ WHERE AMPLITUDE SN = LG
 \$ LG ON PWM AND FCC, NOT IN TRIGGER WINDOW
 \$ SCH, FRB CLOSED
 \$ CBK ASSYMETRIC TRACE; MAX DOWNSWING USED
 \$ HAL LONG CODA, PROBABLY REPRESENTS ATTENUATED LG
 \$ GBN NO LG
 \$ KUQ, IGL, FFC AMP FROM LG
 \$ FFC EARLIER PHASE AT 2141, STONGEST ON SPZ MAY BE BE TELESEISMIC S OR
 \$ UNRELATED TELESEISM
 \$ CLOSE TO LARGE EARTHQUAKE IN 1971
 \$ RELATIVE ARRIVALS ON STJ AND HAL (ONLY COMMON STATIONS) SUGGEST
 \$ SIMILAR LOCATION TO 1971 EARTHQUAKE
 \$

| | | | | | | | | |
|------|-------------|--------|---------|---------|--|---------|----------|-----|
| STJN | 9201130610P | A08074 | D | X0833 | | | | |
| STJN | W 0268KM | 22 | 063 279 | 00 -197 | | 0000000 | 00ML00MN | |
| CBK | 9201130610P | A08575 | | B1033 | | 040 | 135 | 170 |

GBN 9207061658P C6017 C6150 030 317 02 8
 GBN W 0960KM 01 -192 262 49 00 010 0000013 29ML22MN
 Z

92c

+46.118- 47.438F1MN=3.9 0420227 17071992 00.1120.087 0.1 9 17 20.77 218.00 0 1ML=4.2 30 0L3.62
 EASTERN MARGIN OF GRAND BANKS

\$ FIRST EARTHQUAKE KNOWN FROM HERE; APPROX 200 KM SE OF NEAREST CLUSTER
 \$ OF EARTHQUAKES

\$ STJN VERY CLEAR, GOOD PHASES; LG SEEMS OK

\$ CBK GOOD LG

\$ LG FROM THIS SITE TO NEWFOUNDLAND STATIONS IS UNEXPECTED

\$ ALSO THE LG HAS QUITE LONG PERIODS

\$ GBN VERY WEAK LG; LG AT ALL ECTN SITES APPEARS ATTENUATED

\$

\$ ECTN TRIGGER CONTAINS ONLY PN FOR STATIONS BEYOND LMN

\$ KURILE AFTERSHOCK AT THIS TIME INTERFERES: MNQ, MOQ, DAQ READINGS WERE OF THIS EVENT

\$ JAQ DEAD

\$ KUQ LG ARRIVAL NOT READABLE, BUT AMPLITUDE READ AT TIME CORRESPONDING TO LG

\$

| | | | |
|--------------------------------|------------------------------------|--|--|
| STJN 9207170421P A21223 - | A22052 C22185 040 62 127 | | |
| STJN NW 0437KM 31 060 294 49 | 00 002 08 -497 0003218 43ML40MN | | |
| CBK 9207170421P A22122 D | B2332 C2418 070 43 36 | | |
| CBK NW 0852KM 01 -013 295 49 | 12 -152 00 -006 0000751 47ML38MN | | |
| GBN 9207170421P B22407 + | C2424 | | |
| GBN W 1097KM 12 -151 271 49 | 01 -165 0000000 00ML00MN | | |
| HAL 9207170421P C23045 | B25055 050 116 8 8 | | |
| HAL W 1274KM 00 066 268 49 | 24 210 0000087 42ML32MN | | |
| LMN 9207170421P B231104 -1.52 | C251825 0571000 51 8 25243 | | |
| LMN W 1344KM 44 -287 275 49 | 01 -157 0000056 41ML30MN | | |
| SMQ 9207170421P B232985 -0.10 | | | |
| SMQ NW 1500KM 14 -163 295 49 | 0000000 00ML00MN | | |
| GSQ 9207170423P B233080 -0.10 | | | |
| GSQ W 1511KM 21 -197 289 49 | 0000000 00ML00MN | | |
| ICQ 9207170423P XB233268 -0.10 | XC2604 | | |
| ICQ W 1529KM 00 -229 292 49 | 00 636 0000000 00ML00MN | | |
| CNQ 9207170423P XB233972 -0.10 | | | |
| CNQ W 1584KM 00 -203 290 49 | 0000000 00ML00MN | | |
| SLQ 9207170421P C23505 | C2624 X2810 100 68 4 3 | | |
| SLQ W 1649KM 00 099 284 49 | 00 092 00 1181\$ 0000037 42ML30MN | | |
| MNQ 9207170421P -0.10 | C26225 100 96 8 3 | | |
| MNQ NW 1651KM 295 49 | 01 -125 0000052 44ML32MN | | |
| DPQ 9207170421P | X2723 XC2938 050 140 2 3 | | |
| DPQ W 1943KM 281 50 | 00 -258 00 1863\$ 0000018 40ML28MN | | |
| KUQ 9207170421P | 090 44 1 3 | | |
| KUQ NW 1948KM 321 47 | 0000016 40ML28MN | | |

Z

92d

+47.353- 49.114F1MN=3.4 1131520 10081992 00.0830.070 0.0 5 9 10.39 218.00 0 1ML=3.6 30 0L3.62
 GRAND BANKS

\$ IN CLUSTER OF EARTHQUAKES (4 PREVIOUS) NW OF HIBERNIA

\$ VERY POOR AZIMUTHAL COVERAGE FOR THESE EARTHQUAKES

\$ NOT ON DAQ, ICQ

\$ STJ SN>LG

\$ CBK SG=3*SN

\$ GBN AMP READ FROM SN, THERE IS WEAK LG FOLLOWING THE SN

\$ CBK LG MIGHT BE 4 SEC EARLIER

\$

| | |
|-----------------------------|----------------------------------|
| STJ 9208101132P A32312 | B33012 XC3308 040 73 165 8 |
| STJ W 0274KM 00 000 276 49 | 44 129 00 006 0003550 37ML37MN |
| CBK 9208101132P B33211 | B34260 XC3458 040 135 32 |
| CBK W 0682KM 01 017 288 49 | 12 -068 00 -244 0000372 40ML34MN |
| GBN 9208101132P B33572 | B3529 X3627 030 370 12 8 |
| GBN W 0978KM 01 021 262 49 | 09 -059 00 483 0000068 36ML29MN |
| HAL 9208101132P | |
| HAL W 1160KM 260 49 | 0000000 00ML00MN |
| LMN 9208101132P B34255 | C3618 030 428 5 8 |
| LMN W 1213KM 01 -017 268 49 | 04 -162 0000024 34ML26MN |
| MNQ 9208101132P C3455 | XC3933 |
| MNQ W 1479KM 17 -324 291 49 | 00 5223\$ 0000000 00ML00MN |

Z

94a

+52.127- 48.523F1ML=3.2 1052279 08011994 00.0380.065 0.2 6 10 40.56 218.00 0 LMN=0.0 00 0L0.00
NORTHEAST NEWFOUNDLAND SLOPE
\$ LIES WITHIN LINEAR TREND OF EPICENTRES ALONG EXTENSION OF CARTWRIGHT FRACTURE ZONE
\$ SEEN FIRST ON DRLN, PICKED UP WITH DIFFICULTY ON LMN
\$ ANALOG ECTN READ - TRACES ARE MERE WRIGGLES
\$ NOT ON HAL, CBK
\$ KUQ IS WRIGGLE
\$

| | | | | | | | |
|------|-------------|----------------|-------|---------|-------------------|---------|--|
| STJN | 9401081055P | B53455 | | B5444 | 040 62 9 | 8 | |
| STJN | SW 0591KM | 07 -030 213 49 | | 08 033 | 0000228 35ML31MN | | |
| DRLN | 9401081054P | B540047 | +0.00 | B550865 | 0152240 46 0000 8 | 559.62 | |
| DRLN | SW 0710KM | 02 016 247 49 | | 08 -033 | 0000086 30ML28MN | | |
| GBN | 9401081055P | | | C56564 | 030 317 2 | 8 | |
| GBN | SW 1210KM | 237 49 | | 05 105 | 0000013 31ML23MN | | |
| ICQ | 9401081055P | | | C5726 | | | |
| ICQ | W 1350KM | 265 49 | | 05 104 | 0000000 00ML00MN | | |
| LMN | 9401081055P | C552005 | +0.00 | C572900 | 0452560 11 0000 8 | 5736.35 | |
| LMN | SW 1378KM | 12 -163 246 49 | | 18 -195 | 0000006 31ML21MN | | |
| MNQ | 9401081055P | C5529 | | C57415 | | | |
| MNQ | W 1419KM | 23 222.271 49 | | 13 165 | 0000000 00ML00MN | | |
| KUQ | 9401081055P | | | X5738 | | | |
| KUQ | NW 1428KM | 306 49 | | 00 -362 | 0000000 00ML00MN | | |
| FRB | 9401081055P | | | | | | |
| FRB | NW 1743KM | 326 49 | | | 0000000 00ML00MN | | |

Z

94b

+47.827- 52.6730LMN=3.1 1813492 11081994 00.0270.064 0.2 6 11 30.50 2 0.00 0 1ML=3.3 10 0L3.62
ST JOHN'S NFLD ST JOHN'S T.-N.

\$ LOCATION PERHAPS 15-30 KM N OF ST JOHN'S, BUT N-S RESOLUTION
\$ IS POOR BECAUSE OF STATION DISTRIBUTION
\$ BECAUSE OF TIME OF DAY AND PROXIMITY TO CITY A BLAST WAS SUSPECTED
\$ ENQUIRIES BY PAUL BARNES (MEMORIAL UNIV., PAUL@CONVEX.ESD.MUN.CA)
\$ TO A LOCAL EXPLOSIVES SUPPLIER INDICATED:
\$ 1. APPROXIMATELY 20 KM WEST OF ST. JOHN'S (LONG POND), THERE IS A
\$ PYROPHYLITE MINE THAT USES CHARGES ON THE ORDER OF 25 KG.
\$ 2. ON 24 OCTOBER, APPROX 12-15 KM WEST OF ST. JOHN'S, 300X25KG WERE
\$ DETONATED OVER A 2 SECOND INTERVAL TO BLAST AWAY A ROCK FORMATION TO
\$ GENERATE ROADBED GRAVEL. THE MAXIMUM CHARGE SIZE IN ANY ONE HOLE WAS
\$ 250KG. THIS WAS THE LARGEST BLAST EXPLODED. THE DATE MAY BE IN ERROR.
\$ THE EVENT WAS NOT LOCATED ON THE STJN RECORD.
\$ 3. ABOUT 5-8 KM N OF THAT SITE, CONSTRUCTION OF A BYPASS ROAD IS BEING
\$ UNDERTAKEN. ACCORDING TO THE EXPLOSIVES SUPPLIER, THEY ARE DETONATING
\$ CHARGES ON THE ORDER OF 50 KG.
\$ 4. THERE IS ALSO BLASTING FOR ROAD CONSTRUCTION OCCURRING ALONG THE TCH
\$ APPROXIMATELY 50 KM WEST OF ST. JOHN'S.
\$ 5. UNAWARE OF ANY OTHER CONSTRUCTION.
\$
\$ ADAMS CHECKED ALL THE STJN RECORDS HERE FOR AUG 10 TO EARLY OCT AND
\$ IDENTIFIED 10 EVENTS, 3 WITH A S-P OF 5.5 SEC AND AROUND 2030-2230 UT
\$ (PROBABLY SINGLE SOURCE, POSSIBLY YOUR TCH BLASTING?), AND THE REST
\$ WITH S-P AROUND 2 SEC AT 1200-2045 UT (DAYLIGHT HOURS, LIKELY MULTIPLE
\$ SOURCES WITHIN 20 KM). NOT A SINGLE ONE WOULD BE CLASSIFIED AS A LARGE
\$ BLAST, AND THE LARGEST EVENT IS AT LEAST 1 MAGNITUDE BELOW THE AUG 11TH
\$ EVENT. IF THE AUG 11 EVENT WAS A BLAST, WE SEEM FORCED TO CONCLUDE THAT
\$ IT WOULD HAVE NEEDED TO INVOLVE C. 10* THE 300X25 KG = 87 TONNES, AND
\$ THEREFORE WOULD HAVE BEEN REMEMBERED!
\$ ---> CONCLUSION CLASSIFY AS AN EARTHQUAKE
\$
\$ MISSED DURING REGULAR DAILY PROCESSING OF CSN
\$ STJN CODA LASTS 2.5 MIN
\$ STJN: SUSPECT S-P IS ABOUT 2.8 SEC, BUT UNREADABLE
\$ HAL NOT EVIDENT ON NOISY RECORD
\$ SCH NOISY RECORD
\$ LMN DEAD
\$ MNQ FROM ANALOG
\$ DONE BY ADAMS
\$ LAST COMPARABLE EVENT WAS IN 1969:
\$ +47.632- 52.156F1MLN=3.4 2153207 05081969 00.0860.110 0.2 5 9 40.94 218.00
\$ THIS WAS FROM ALMOST DUE EAST OF STJ ACCORDING TO 3 COMPONENT MOTIONS
\$

| | | | | | | | |
|------|-------------|----------------|-------|-----------------|--------------------|---------|--|
| STJN | 9408111814P | | | A13542 C | | | |
| STJN | S 0029KM | 192-90 | | 06 024 | 0000000 00ML00MN | | |
| DRLN | 9408111814P | A144492C | -1.00 | A152635 B153947 | 0252303 672 0000 0 | 1544.58 | |
| DRLN | NW 0391KM | 32 -055 296 49 | | 11 033 11 129 | 0000733 33ML33MN | | |

| | | | |
|----------------------------|--------|-----------------------------------|---------|
| GBN 9408111814P A15257 | | B16355 B1709 040 216 20 | 3 |
| GBN W 0729KM 00 001 252 49 | | 14 -144 15 -151 0000145 37ML30MN | |
| MNQ 9408111818P | | C1920 060 164 10 | |
| MNQ W 1210KM | 290 49 | 05 -356 0000064 41ML30MN | |
| SCH 9408111818P | -1.00 | B183037 C193952 0452560 87 0000 0 | 1942.17 |
| SCH NW 1251KM | 314 49 | 11 130 05 362 0000047 39ML29MN | |
| LMQ 9408111818P | -1.00 | C184462 | 0 |
| LMQ W 1324KM | 275 49 | 00 010 0000000 00ML00MN | |

Z

94c

+47.315- 51.72501MN=2.6 0932259 01121994 00.1150.138 0.4 5 10 30.94 218.00 0 1ML=2.5 30 0L3.62
OFFSHORE NEWFOUNDLAND AU LARGE DE TERRE-NEUVE
\$ VERY CLEAR ON ANALOG STJN RECORD; ALMOST IGNORED ON DRLN ANALOG
\$ LOCATION USING DAN IMPOSSIBLE, KEPT CRASHING; READINGS WERE TYPED IN MANUALLY
\$ POOR LOCATION RESOLUTION IN N-S DIRECTION
\$ RESIDUALS ARE LOWER FOR A SHALLOW DEPTH
\$ NO LG ON GBN OR LMN
\$ NOT ON HAL
\$ NOT REPORTED FELT
\$ PREVIOUS CLOSE EVENTS ARE
\$ +47.632- 52.156F1MN=3.4 2153207 05081969 00.0860.110 0.2 5 9 40.94 218.00
\$ +47.827- 52.67301MN=3.1 1813492 11081994 00.0270.064 0.2 6 11 30.50 2 0.00
\$ DONE BY ADAMS

| | | | | |
|------------------------------|--------|----------|----------------------------------|---------|
| STJN 9412010932P | | A32387 D | B3250 025 120 23 | |
| STJN W 0083KM | 290-78 | 37 -087 | 01 068 0004817 26ML30MN | |
| DRLN 9412010932P A33307 | | | B34204 C34386 0357138 428 0000 | 3448.58 |
| DRLN NW 0480KM 15 055 299 49 | | | 24 283 00 -008 0000108 29ML26MN | |
| GBN 9412010932P B34070 | | | B35245 020 563 1 | 8 |
| GBN W 0782KM 00 -002 258 49 | | | 20 261 0000006 20ML17MN | |
| LMN 9412010932P B34344 | | | B36113 | |
| LMN W 1015KM 03 -095 266 49 | | | 00 -003 0000000 00ML00MN | |
| SCH 9412010932P C35161 | | | XC37283 X38414 0354999 25 0000 0 | 3854.75 |
| SCH NW 1342KM 00 087 314 49 | | | 00 740 00 469 0000009 32ML22MN | |

Z

95

+47.840- 52.517F1MN=2.4 0646205 22011995 00.0180.049 0.0 3 7 10.29 218.00 0 1ML=2.2 20 0L3.62
33 KM NE OF ST JOHN'S, NFLD 33 KM NE DE ST JOHN'S, T.-N.
\$ STJN VERY CLEAR, WRONG TIME OF DAY FOR BLAST
\$ LMN DEAD
\$ GBN TINY SIGNAL, SN STRONGER THAN PN
\$ DRLN CLEAR WHEN FILTERED
\$ DRLN READINGS FROM VERTICAL COMPONENT ONLY, ONE HORIZONTAL HAD C. 30 SEC TIMING ERROR
\$ TIMING DUBIOUS ON THE RECORDS AS TRANSFORMED USING RDSEED AND READ
\$ SCHQ TRIED BUT NOT VISIBLE THERE
\$ NOT ON HAL
\$ LAST NEARBY EVENT 940811 M3.1, MIGHT BE IN SAME PLACE
\$ DONE BY ADAMS

| | | | | |
|------------------------------|--------|----------|----------------------------------|--------|
| STJN 9501220647P | | A46269 D | A46315 020 146 195 | |
| STJN SW 0035KM | 211-63 | 01 -005 | 00 002 0004196 19ML31MN | |
| DRLN 9501220647P C47196 | -1.00 | | B47569 B48117 0207211 237 0000 0 | 4813.5 |
| DRLN NW 0401KM 18 347 295 49 | | | 08 058 11 -069 0000103 24ML24MN | |
| GBN 9501220647P C48012 | | | B49077 | |
| GBN W 0740KM 32 465 252 49 | | | 00 013 0000000 00ML00MN | |
| SCHQ 9501220654P | +0.00 | | XX545490 | 0 |
| SCHQ NW 1262KM | 313 49 | | 00 9999\$ 0000000 00ML00MN | |

Z

[SAMDEV.IOC]PICKF
C. WONG MAY 1

PIK FILE

>>>>>>> NEW FORMAT VERSION AUGUST 1987
>>>>>>> 5 CHARACTER STATION ID

The PIK file is the input file to and also the output file (o version newer) from the CANSESS MULTILAYER epicenter location prog (IOC). SAM PIK (or PK4) command generates a PIK file automatically the event. These PIK files can be modified/created by the EPK progr or by the DEC text editor EDT.

It contains four types of records:

1. ESR - earthquake solution record.
2. ECR - earthquake comment record.
3. ODR - observed data record.
4. CDR - calculated data record.

The ESR (if it exists) has to be the first record in the file. It is an output record from the previous IOC. Some of its fields can be modified by the IOC parameters for the next run. The ECR records come the ESR, they have to be before the first ODR. There is one ODR for station with picked information. This is an input record, IOC progr not modified any field on this record. The CDR contains only the ca results for this station. If it exists, it is always right after it corresponding ODR. You can have as many ECR, ODR or sets of ODR & C you want. However, the current EPK program can handle total of 100 at most, and the IOC program is dimensioned only 100 for all the ph picked. If you have any difficulties with the above limitations, ple me know. The detail layout of these records starts on the next pag

EARTHQUAKE SOLUTION RECORD (ESR)

(solution record is defined as having "I" or "-" on col.1 and "M" ON

| CLOUNS | ENTRY | FORMAT | DEFINITION |
|--------|--------------|--------|--|
| 1-1 | + | A1 | PRIME SOLUTION BY EPB |
| | - | | PRIME SOLUTION BY OTHER AGENCY |
| | 0 | | SUPPLEMENTARY SOLUTION |
| 2-7 | 45.233 | F6.3 | NORTH LATITUDE, DEGREES |
| 8-15 | -123.300F8.3 | | LONGITUDE, DEGREES |
| 16-16 | O | A1 | HYPOCENTRE QUALITY INDICATOR. |
| | F | | POOR QUALITY SOLUTION |
| | | | GOOD QUALITY SOLUTION |
| 17-17 | BLANK | I1(A1) | OBSERVED DATA FORMAT INDICATOR, PRE-1979 DATA FORMAT USED |
| | 1 | | 1979 DATA FORMAT USED. |
| 18-19 | ML | A2 | PRIME MAGNITUDE TYPE |
| | MIE | | RICHTER |
| | MN | | EBEL |
| | MLG | | NUTTLI (DEFAULT) |
| | MB | | H. & K. |
| | MS | | BODY-WAVE |
| | MC | | SURFACE WAVE |
| | | | CODA LENGTH |
| 20-20 | BLANK | | |
| 21-23 | 3.1 | F3.1 | AVERAGE MAGNITUDE VALUE (MM) |
| 24-24 | BLANK | | |
| 25-26 | 18 | I2 | ORIGIN TIME HOUR, U.T. |
| 27-28 | 23 | I2 | ORIGIN TIME MINUTE |
| 29-31 | 323 | I3 | ORIGIN TIME SECOND*10 (OR F3.1 IN SECOND) |
| 32-32 | BLANK | | |
| 33-34 | 12 | I2 | DAY |
| 35-36 | 03 | I2 | MONTH |
| 37-40 | 1979 | I4 | YEAR |
| 41-41 | BLANK | | |
| 42-42 | 2 | I1 | STANDARD DEVIATION ORIGIN TIME, SECONDS |

| | | | |
|-------|-------|------|---|
| 43-47 | 0.122 | F5.3 | STD IN LATITUDE, DEGREES |
| 48-52 | 0.333 | F5.3 | STD IN LONGITUDE, DEGREES |
| 53-53 | BLANK | | |
| 54-56 | 0.3 | F3.1 | STD IN MAGNITUDE FOR EPB |
| | XXX | A3 | AGENCY CODE FOR EXTERNAL MAG, DEPENDS ON COL. |
| 57-59 | 34 | I3 | NUMBER OF STATIONS USED FOR HYPOCENTER |
| 60-62 | 14 | I3 | NUMBER OF PHASES USED FOR THIS HYPOCENTER. |
| 63-65 | 14 | I3 | NUMBER OF AMPLITUDE USED FOR MAGNITUDE. |
| 66-69 | 0.33 | F4.2 | RMS OF HYPOCENTER SOLUTION, SECONDS. |
| 70-70 | BLANK | A1 | SOLUTION TYPE INDICATOR |
| | Z | | FIXED DEPTH. |
| | X | | FREE DEPTH |
| | N | | NO ACTION FOR THE WHOLE FILE |
| | N | | ASSIGNED HYPOCENTER AND TIME |
| | H | | ASSIGNED HYPOCENTER, BUT CALCULATED ORIGIN TIME. |
| 71-71 | | I1 | AGENCY CODE |
| | 1 | | USGS |
| | 2 | | EPB |
| | 3 | | PGC |
| | 4 | | SEA |
| | 5 | | NEIS |
| | 6 | | ISC |
| | 7 | | LDGO |
| | 8 | | WES |
| | 9 | | UAGI |
| | | | UNIVERSITY OF WASHINGTON |
| | | | NATIONAL EARTHQUAKE INFORMATION CENTER |
| | | | INTERNATIONAL SEISMOLOGICAL CENTER |
| | | | LAWOMT-DOHERTY GEOLOGICAL OBSERVATORY |
| | | | WESTON GEOPHYSICAL OBSERVATORY |
| | | | UNIV. OF ALASKA, GEOPHYSICAL INSTITUTE |
| 72-76 | 18.33 | F5.2 | FOCAL DEPTH, KM |
| | | | IF AND ONLY IF COL. =Z, FREE DEPTH SOLUTION |
| 77-80 | 3.0 | F4.1 | STD IN DEPTH, KM (OR I4 FORMAT IN 100-METERS) |
| 81-82 | 03 | I2 | MODEL NUMBER |
| 83-85 | ML# | | AVERAGE ML |
| 86-88 | 1.3 | F3.1 | RICHTER MAGNITUDE |
| 89-91 | 008 | I3 | NUMBER OF STATIONS USED TO CALCULATE AVERAGE M |
| 92-92 | 1 | I1 | MULTILAYER HYPO SIMULATION FLAG, 0-OFF, 1-ON. |
| 93-93 | F | A1 | FELT |
| | N," " | | NOT FELT |
| | | | ! FLOO1 |
| | | | ! FLOO1 |
| | | | ! FLOO1 |
| 94-95 | 10 | I2 | NB. OF ASSOCIATED EVENTS |
| | | | ! FLOO1 |
| 96-96 | E," " | A1 | EARTHQUAKE |
| | B | | BLAST |

R
 P
 97-100 3.56 F4.2 S VELOCITY USED BY SINGLE LAYER MODEL ! FLOO
 ROCKBURST
 POSSIBLE ROCKBURST
 < FORMAT(A1,F6.3,F8.3,2A1,A2,1X,F3.1,1X,I2.2,I2.2,I3.3,1X,2I2.2,
 & '19',I2.2,1X,I1,2F5.3,1X,A3, ! FOR AG
 & '19',I2.2,1X,I1,2F5.3,1X,F3.1, ! FOR ST
 & 3I3.3,F4.2,A1,I1,F5.2,I4,T81,I2,'MC-',F3.1,I3,I1
 & A1,I2,A1,F4.2)

EARTHQUAKE COMMENT CARDS (ECR)

| COLUMNS | ENTRY | FORMAT | DEFINITION |
|---------|-------|--------|-----------------------------------|
| 1-40 | 40A1 | | EARTHQUAKE DESCRIPTION IN ENGLISH |
| 41-80 | 40A1 | | EARTHQUAKE DESCRIPTION IN FRENCH |

OBSERVED DATA RECORD (ODR)

| COLUMNS | ENTRY | FORMAT | DEFINITION |
|---------|---------|--------|---|
| 1-5 | OTF | A5 | STATION CODE ! FL001 |
| 6-7 | 79 | I2 | YEAR ! FL001 |
| 8-9 | 12 | I2 | MONTH ! FL001 |
| 10-11 | 23 | I2 | DAY ! FL001 |
| 12-13 | 12 | I2 | HOURL, U. T. ! FL001 |
| 14-15 | 14 | I2 | MINUTE OF 1ST RECORDED P PHASE, NOT NECESSARY @ |
| 16-16 | P | A1 | INSTRUMENT CODE ! FL001 |
| | L | | SHORT PERIOD INSTRUMENT READ |
| | | | LONG PERIOD INSTRUMENT READ, AMP. & 1ST MOTION |
| 17-17 | BLANK | | ! FLO |
| 18-18 | " " | A1 | PN WEIGHT ! FLO |
| | X | | USED IN CALCULATION |
| | | | NOT USED IN CALCULATION |
| 19-19 | A | A1 | PN QUALITY DESIGNATOR ! FLO |
| | B, " | | SHARP CLEAR BEGINNING (+- 0.25 SEC.) |
| | C | | GOOD BEGINNING (+- 1.0 SEC.) |
| | X | | WEAK POOR BEGINNING (+- 4.0 SEC. OR MORE) |
| | 0 | | PHASE NOT USED IN SOLUTION, LARGE RESIDUAL. |
| 20-21 | 14 | I2 | MINUTE OF PN ARRIVAL ! FLO |
| 22-25 | 2341 | F4.2 | SECOND OF PN ARRIVAL ! FLO |
| 26-28 | CNM ?? | 3A1 | FIRST MOTION OF PN ARRIVAL ! FLO |
| 29-33 | +0.03 | F5.0?? | TIME CORRECTION ! FLO |
| 34-34 | " " | A1 | PG WEIGHT ! FLO |
| | X | | USED IN CALCULATION |
| | | | NOT USED IN CALCULATION |
| 35-35 | A, B, . | A1 | PG QUALITY DESIGNATOR, SEE 16 ! FLO |
| 36-37 | 14 | I2 | MINUTE OF PG ARRIVAL ! FLO |
| 38-41 | 264 | F4.2 | SECOND OF PG ARRIVAL ! FLO |
| 42-44 | DSE | 3A1 | FIRST MOTION OF PG ARRIVAL ! FLO |
| 45-45 | | A1 | SN WEIGHT ! FLO |

USED IN CALCULATION
NOT USED IN CALCULATION

46-46 A,B... A1 SN QUALITY DESIGNATOR, SEE 16 ! FLO

47-48 14 I2 MINUTE OF SN ARRIVAL ! FLO

49-52 52 F4.2 SECOND OF SN ARRIVAL ! FLO

53-53 " " A1 LG WEIGHT ! FLO

54-54 A,B... A1 LG QUALITY DESIGNATOR, SEE 16 ! FLO

55-56 14 I2 MINUTE OF LG ARRIVAL ! FLO

57-60 589 F4.2 SECOND OF LG ARRIVAL ! FLO

61-61 BLANK ! FLO

62-64 031 F3.2 PERIOD OF MAX. TRACE AMPLITUDE, SECOND ! FLO

65-68 150 F4.0 MAGNIFICATION OF INSTRUMENT AT GIVEN PERIOD, I

69-72 125 F4.1 TRACE AMPLITUDE(ONE-HALF MAX. PEAK-TO-PEAK) I ! FLO

73-73 BLANK ! FLO

74-77 I4 DURATION IN SECONDS. ! FLO

78-79 BLANK ! FLO

80-80 BLANK ! FLO

1 MAGNITUDE CODE

2 AMPLITUDE SUITABLE FOR NUTTLI OR RICHTER SCALE

3 AMPLITUDE SUITABLE FOR RICHTER ONLY, CORDILLER

4 AMPLITUDE SUITABLE FOR EHEL

5 AMPLITUDE UNRELIABLE, NOT USED FOR MAGNITUDE

6 AMPLITUDE SUITABLE FOR HUEM & KISCO

8 AMPLITUDE SUITABLE FOR MS SCALE ONLY

SN AMPLITUDE READ,
USE RICHTER SCALE ONLY BEYOND 600 KM IF REQUIR

81-83 BLANK ! FLO

84-85 15 I2 MINUTE OF THE MAX. AMPLITUDE ! FLO

86-89 155 F4.2 SECONDS OF THE MAX. AMPLITUDE ! FLO

< FORMAT(A5,5I2,A1,1X,2A1,I2,F4.2,A3,F5.0,2A1,I2,F4.2,A3,2A1

< 12,F4.2,2A1,I2,F4.2,1X,F3.2,F4.0,F4.1,1X,I4,2X,I

< 3X,I2,F4,2)

CALCULATED DATA RECORD (CDR)

| COLUMNS | ENTRY | FORMAT | DEFINITION |
|---------|-------------|--------|--|
| 1-5 | OTT | A3 | STATION CODE ! FLO01 |
| 6-6 | BLANK | | |
| 7-8 | NW | A2 | QUADRANT OF STATION |
| 9-9 | BLANK | | |
| 10-13 | 1305 | I4 | EPICENTRAL DISTANCE, KM |
| 14-15 | KM | A2 | RECORD FLAG |
| 16-16 | BLANK | | |
| 17-18 | 28 | F2.1 | PN WEIGHT USED FOR CALCULATIONS ! FLO01 |
| 19-19 | BLANK | | |
| 20-23 | 0107 | F4.2 | PN RESIDUAL, SECOND ! FLO01 |
| 24-24 | BLANK, # A1 | | LARGE RESIDUAL FLAG ! FLO01 |
| 25-27 | 235 | I3 | AZIMUTH TO STATION, DEGREES |
| 28-30 | 049 | I3 | EMERGENT ANGLE PN POSITIVE PG NEGATIVE |
| 31-34 | BLANKS | | |
| 35-36 | 14 | F2.1 | PG WEIGHT ! FLO01 |
| 37-37 | BLANK | | |
| 38-41 | -091 | F4.2 | PG RESIDUAL, SECOND ! FLO01 |
| 42-42 | BLANK, # A1 | | LARGE RESIDUAL FLAG ! FLO01 |
| 43-45 | BLANKS | | |
| 46-47 | 07 | F2.1 | SN WEIGHT ! FLO01 |
| 48-48 | BLANK | | |
| 49-52 | 0024 | F4.2 | SN RESIDUAL, SECOND ! FLO01 |
| 53-53 | BLANK, # A1 | | LARGE RESIDUAL FLAG ! FLO01 |

54-55 07 F2.1 SG WEIGHT ; FL001
56-56 BLANK
57-60 -434 F4.2 SG RESIDUAL, SECOND ; FL001
61-61 BLANK, # A1 LARGE RESIDUAL FLAG ; FL001
62-63 BLANKS
64-70 0001356 I7 GROUND VELOCITY, NM/SEC
71-71 BLANK
72-73 35 F2.1 RICHTER OR SURFACE WAVE MAGNITUDE
74-75 ML,MS A2 MAGNITUDE DESIGNATOR
76-77 34 F2.1 NUTTLI MAGNITUDE
78-79 MN A2 MAGNITUDE DESIGNATOR

< FORMAT(A5,1X,A2,1X,14.4,'KM',1X,F2.1,1X,F4.2,A1,2I3.3, >
< & 4X,F2.1,1X,F4.2,A1,3X,F2.1,1X,F4.2,A1,F2.1,1X,
< & F4.2,A1,2X,I7.7,1X,2(F2.1,A2)) >